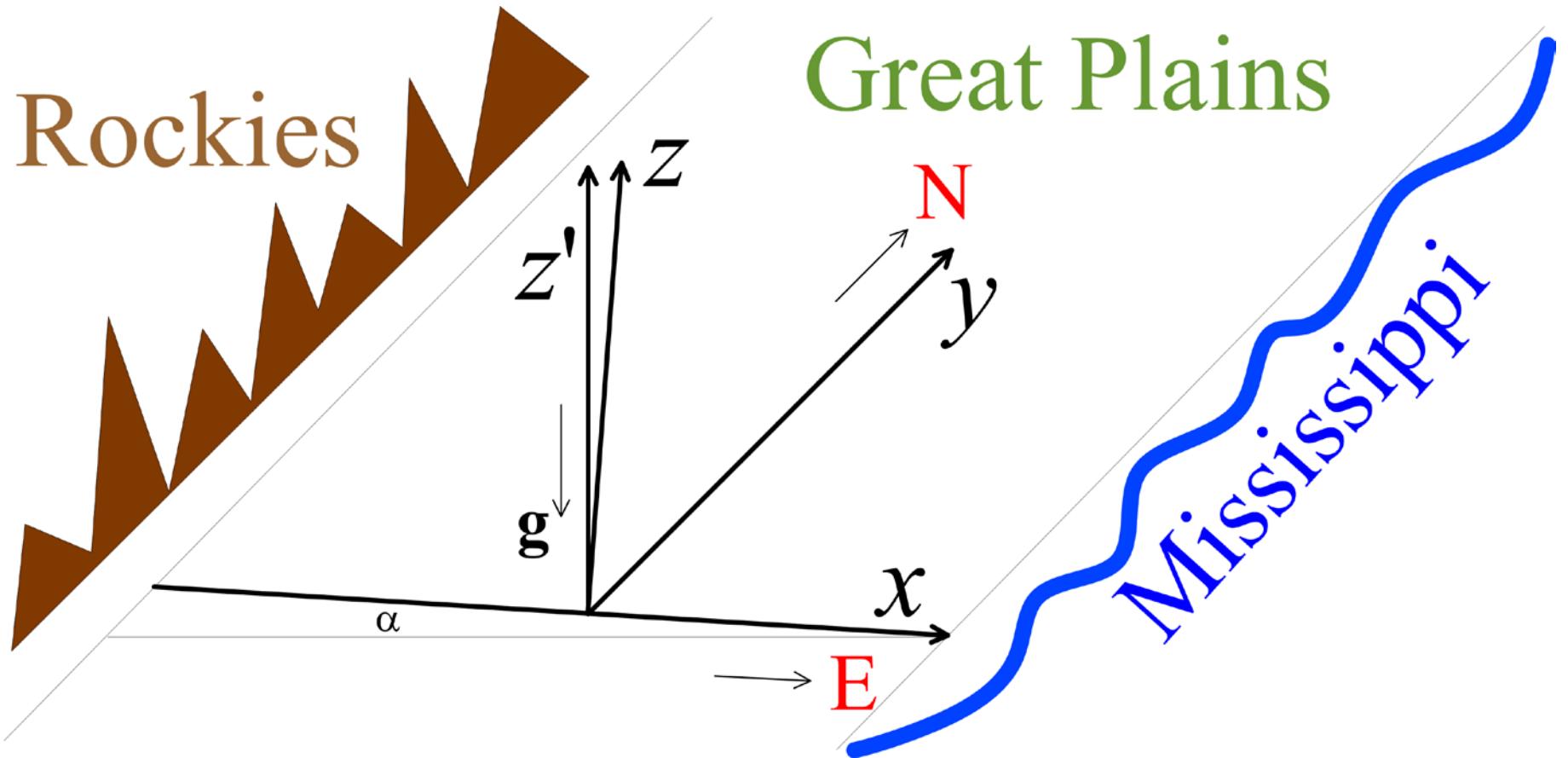


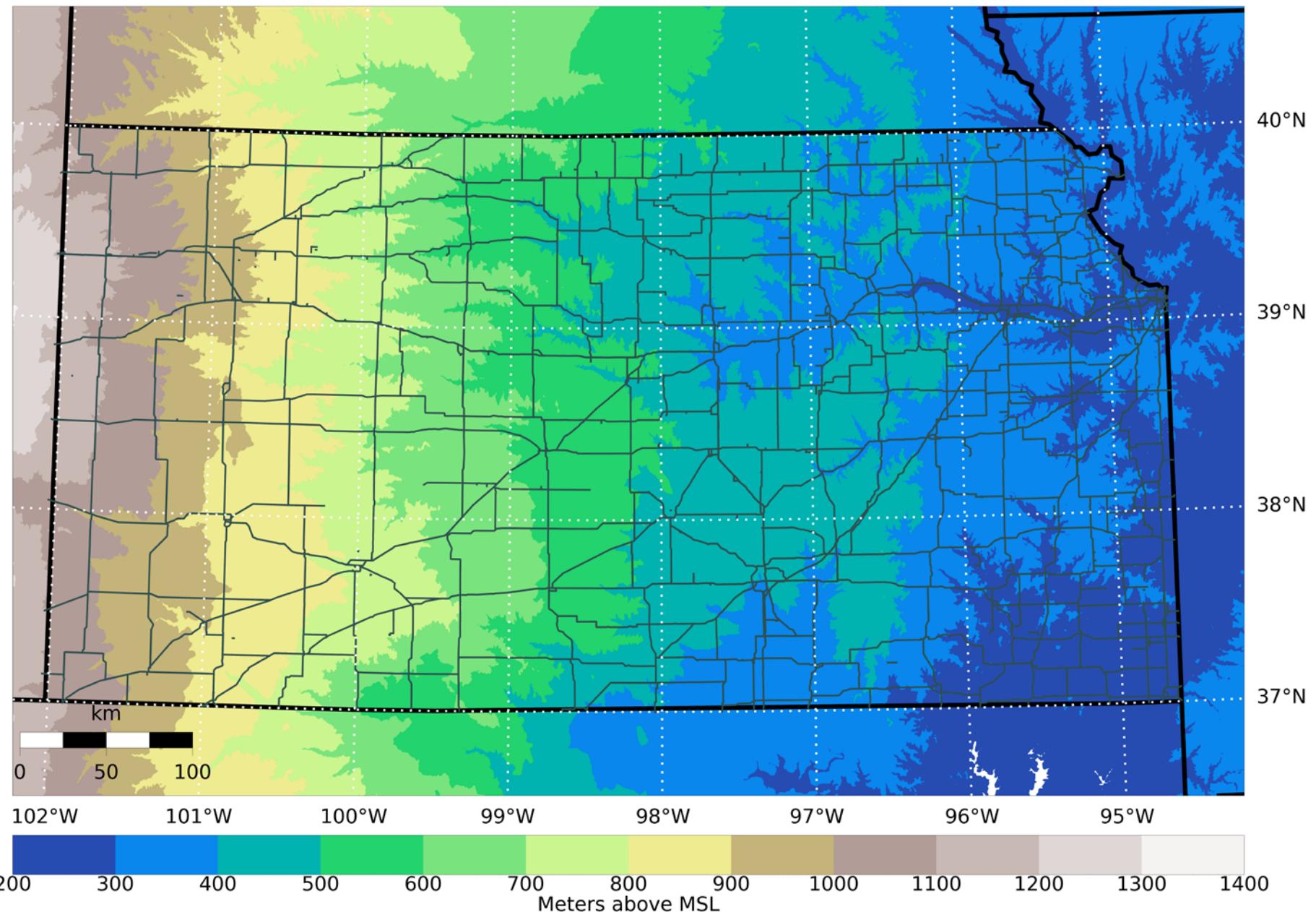
# Effects of shallow slope on the evolution of numerically simulated nocturnal low-level jets

Evgeni Fedorovich, Jeremy Gibbs, and Alan Shapiro

*University of Oklahoma and University of Utah*

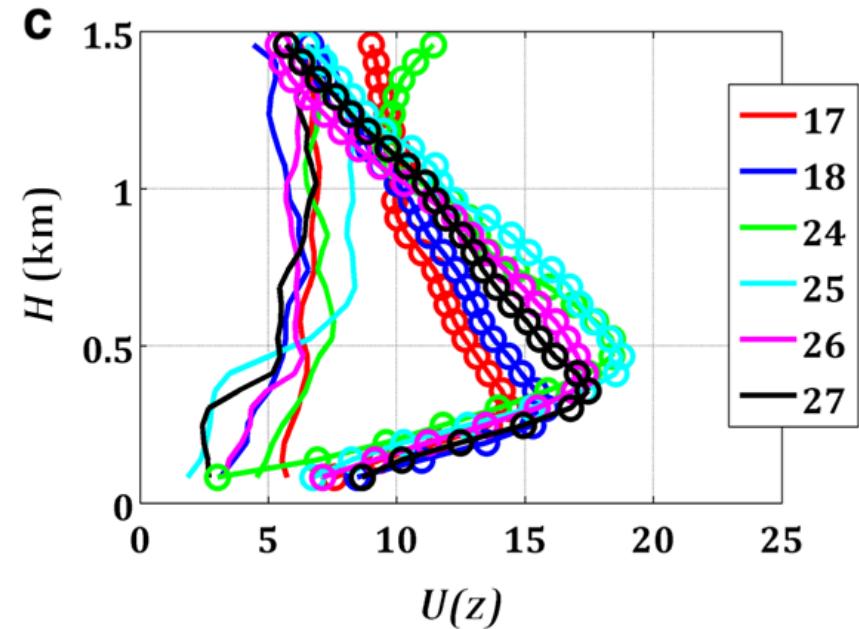
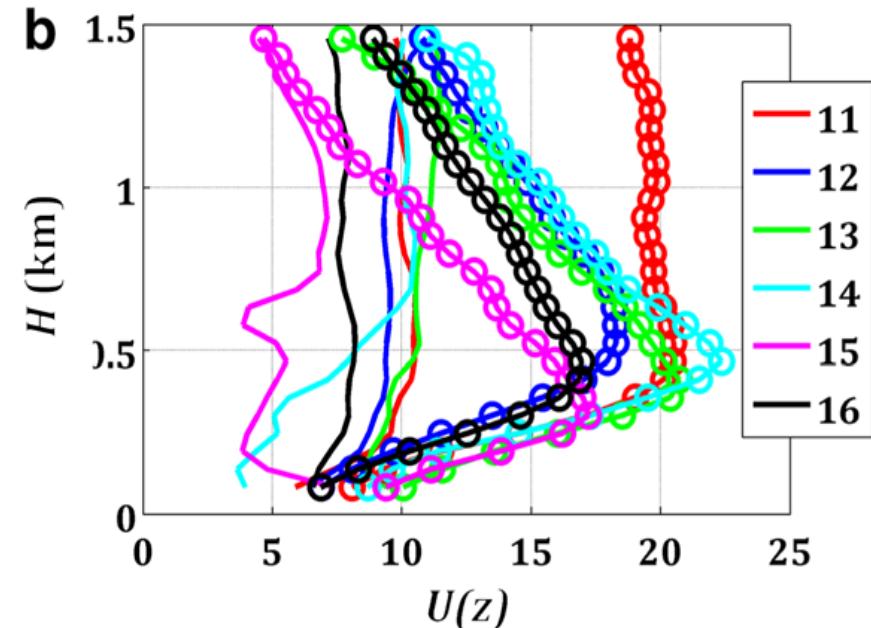
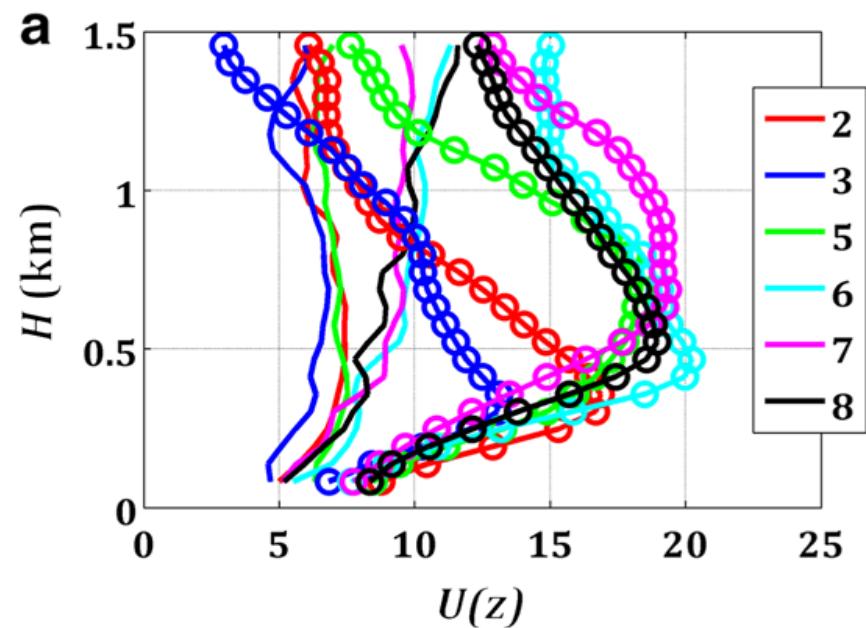


# Terrain of the U.S. Great Plains

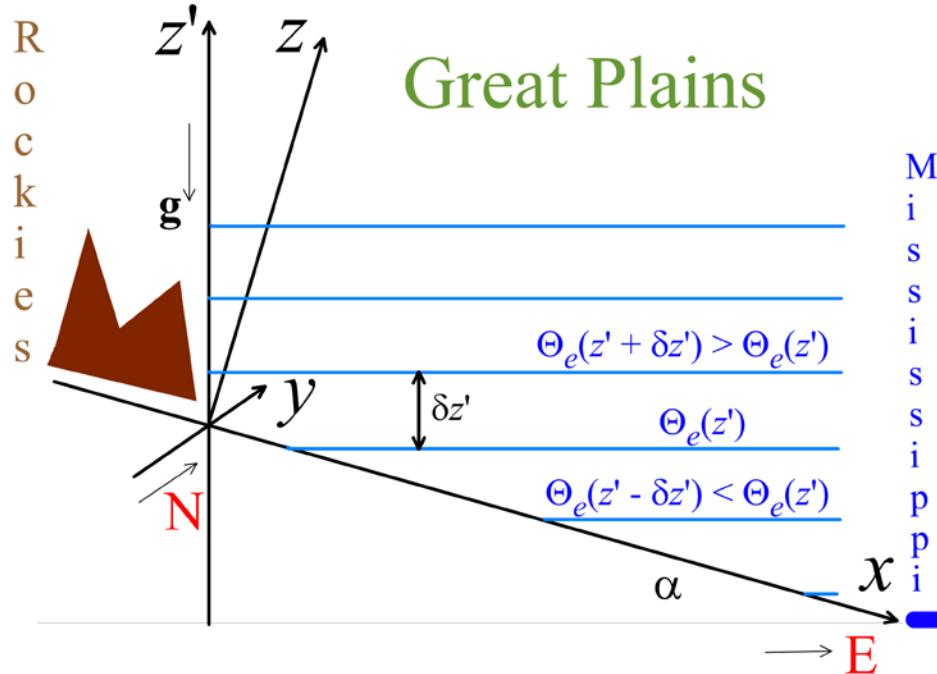


# Nocturnal low-level jets over U.S. Great Plains

(examples from Klein et al. *BLM* 2016)



# Flow variables in slope-following coordinates



$u, v, w$  – velocity components along slope-following coordinates  $x, y, z$ ;

$b = g[\Theta - \Theta_e(z')]/\Theta_r$  – buoyancy;  $\Theta_e(z')$  – environmental potential temperature;

$\pi = [p - p_e(z')]/\rho_r$  – normalized  $p$ ;  $\gamma = d\Theta_e / dz' > 0$  – prescribed  $\Theta_e$  gradient;

$f > 0$  – Coriolis parameter,  $V_g$  – geostrophic wind (known and directed along  $y$ );

$N = (g\gamma/\Theta_r)^{1/2}$  – Brunt-Väisälä frequency;  $\alpha$  – slope angle ( $\sim \pi/1000 = 0.18^\circ$ );

$\nu$  – kinematic viscosity,  $\nu_h$  – thermal diffusivity (we take  $\nu_h = \nu$ , so  $\text{Pr} = \nu/\nu_h$  is 1).

# Governing equations for boundary-layer flow on a slope

Momentum balance in Boussinesq approximation:

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} + w \frac{\partial u}{\partial z} = - \frac{\partial \pi}{\partial x} + f(v - V_g) - b \sin \alpha + \nu \left( \frac{\partial^2 u}{\partial x^2} + \frac{\partial^2 u}{\partial y^2} + \frac{\partial^2 u}{\partial z^2} \right), \quad (1)$$

$$\frac{\partial v}{\partial t} + u \frac{\partial v}{\partial x} + v \frac{\partial v}{\partial y} + w \frac{\partial v}{\partial z} = - \frac{\partial \pi}{\partial y} - fu + \nu \left( \frac{\partial^2 v}{\partial x^2} + \frac{\partial^2 v}{\partial y^2} + \frac{\partial^2 v}{\partial z^2} \right), \quad (2)$$

$$\frac{\partial w}{\partial t} + u \frac{\partial w}{\partial x} + v \frac{\partial w}{\partial y} + w \frac{\partial w}{\partial z} = - \frac{\partial \pi}{\partial z} + b \cos \alpha + \nu \left( \frac{\partial^2 w}{\partial x^2} + \frac{\partial^2 w}{\partial y^2} + \frac{\partial^2 w}{\partial z^2} \right), \quad (3)$$

Thermal energy (buoyancy balance):

$$\frac{\partial b}{\partial t} + u \frac{\partial b}{\partial x} + v \frac{\partial b}{\partial y} + w \frac{\partial b}{\partial z} = N^2(u \sin \alpha - w \cos \alpha) + \nu_h \left( \frac{\partial^2 b}{\partial x^2} + \frac{\partial^2 b}{\partial y^2} + \frac{\partial^2 b}{\partial z^2} \right), \quad (4)$$

Mass conservation (continuity):

$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0. \quad (5)$$

# Simulation setup

## Boundary conditions

**Lateral** ( $x$ - $y$ ) conditions for  $u$ ,  $v$ ,  $w$ ,  $b$ , and  $\pi$  are periodic.

**Upper** conditions (large  $z$ ):  $\partial\varphi / \partial z = 0$ , where  $\varphi = (u, v, b)$ ;  $w = 0$ , and  $\partial\pi / \partial z$  from the third equation of motion (3).

**Lower** conditions ( $z = 0$ ): no-slip/impermeability ( $u = v = w = 0$ ),  $\partial\pi / \partial z$  from (3), and  $b = b_s(t)$  or  $B = -v(\partial b / \partial z) \Big|_{z=0} = B_s(t)$ .

## Initial state

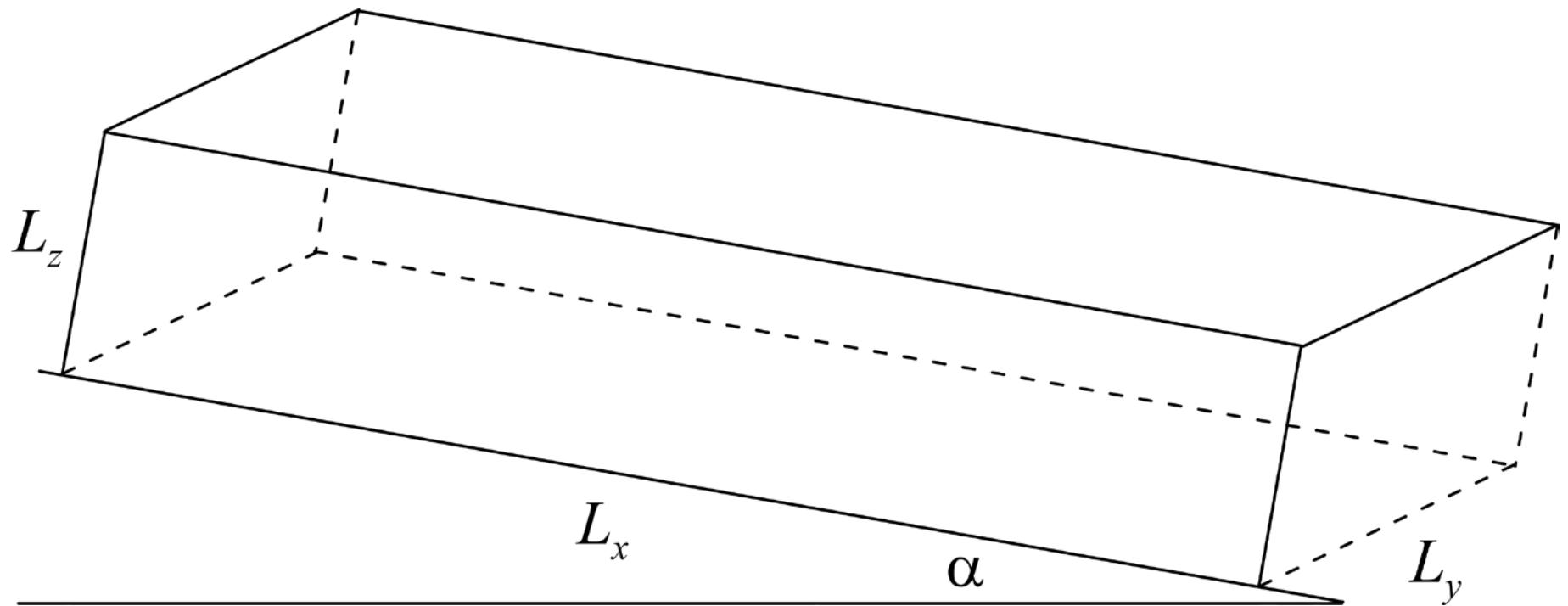
Stably stratified atmosphere in the geostrophic equilibrium with linear  $\Theta_e(z')$  and zero  $b$ .

## Surface buoyancy forcing

At  $0 < t \leq \Delta t_d$ :  $b_{sd} = const > 0$       or       $B_{sd} = const > 0$ .

At  $t > \Delta t_d$ :  $b_{sn} = const < b_{sd}$       or       $B_{sn} = const < 0$ .

# Simulation domain and grid



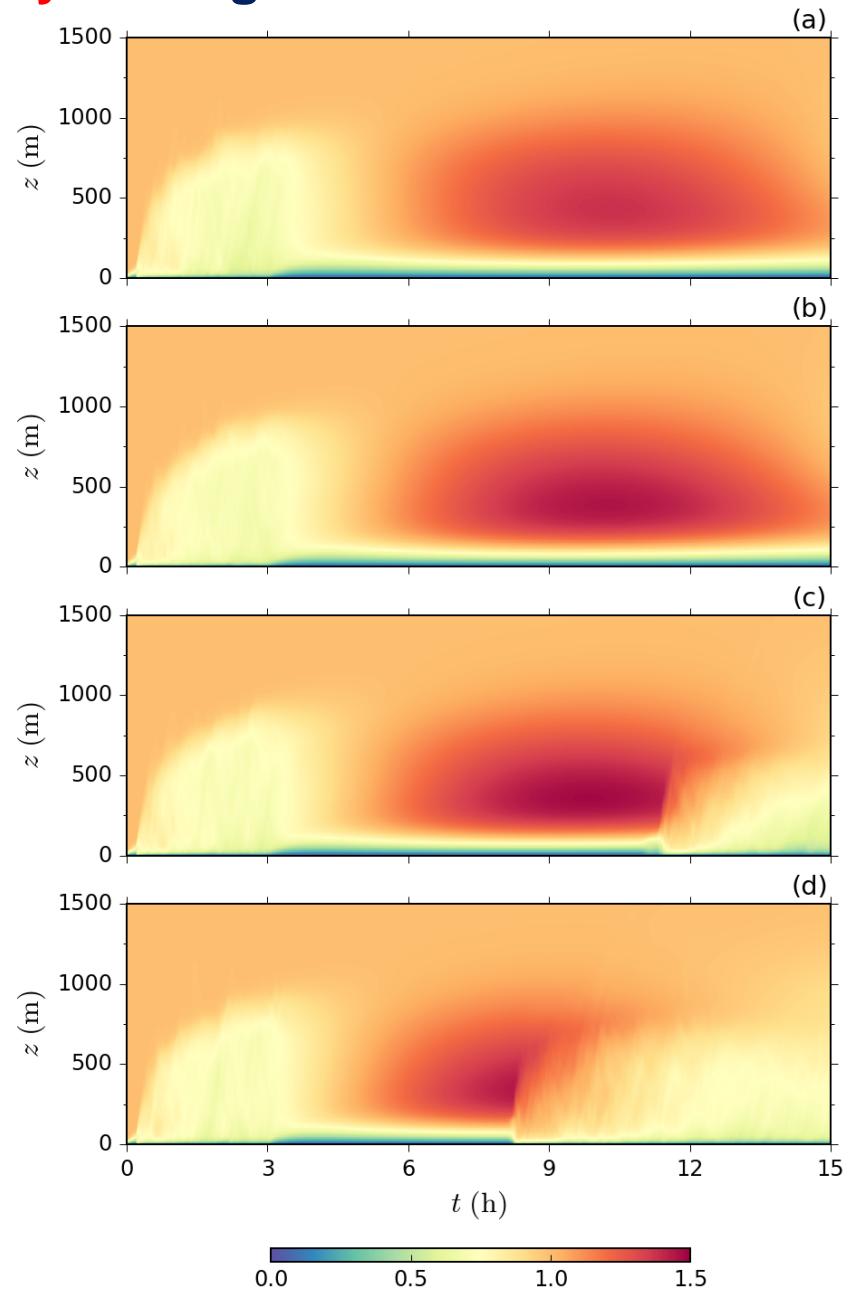
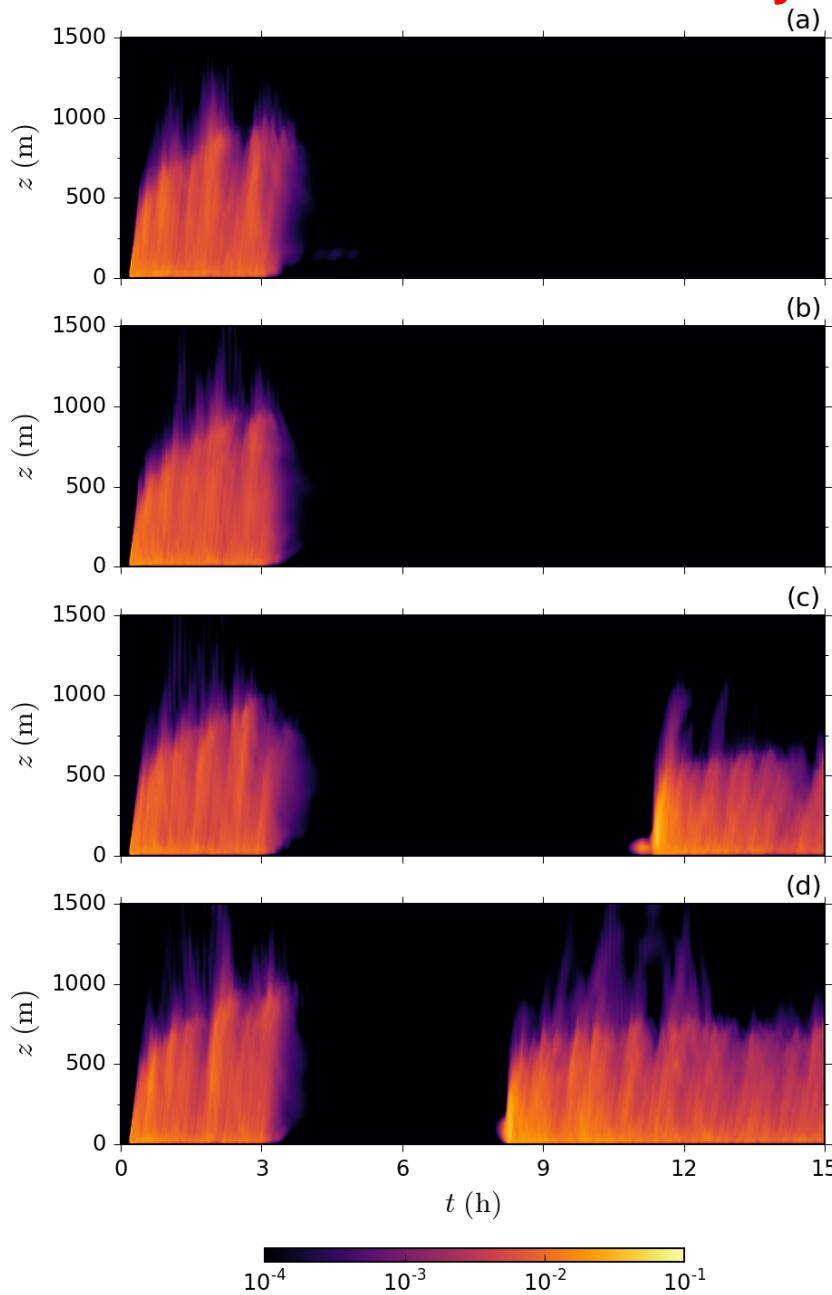
**Dimensions ( $L_i$ , rescaled to VAT) and grid node numbers  $N_i$ :**

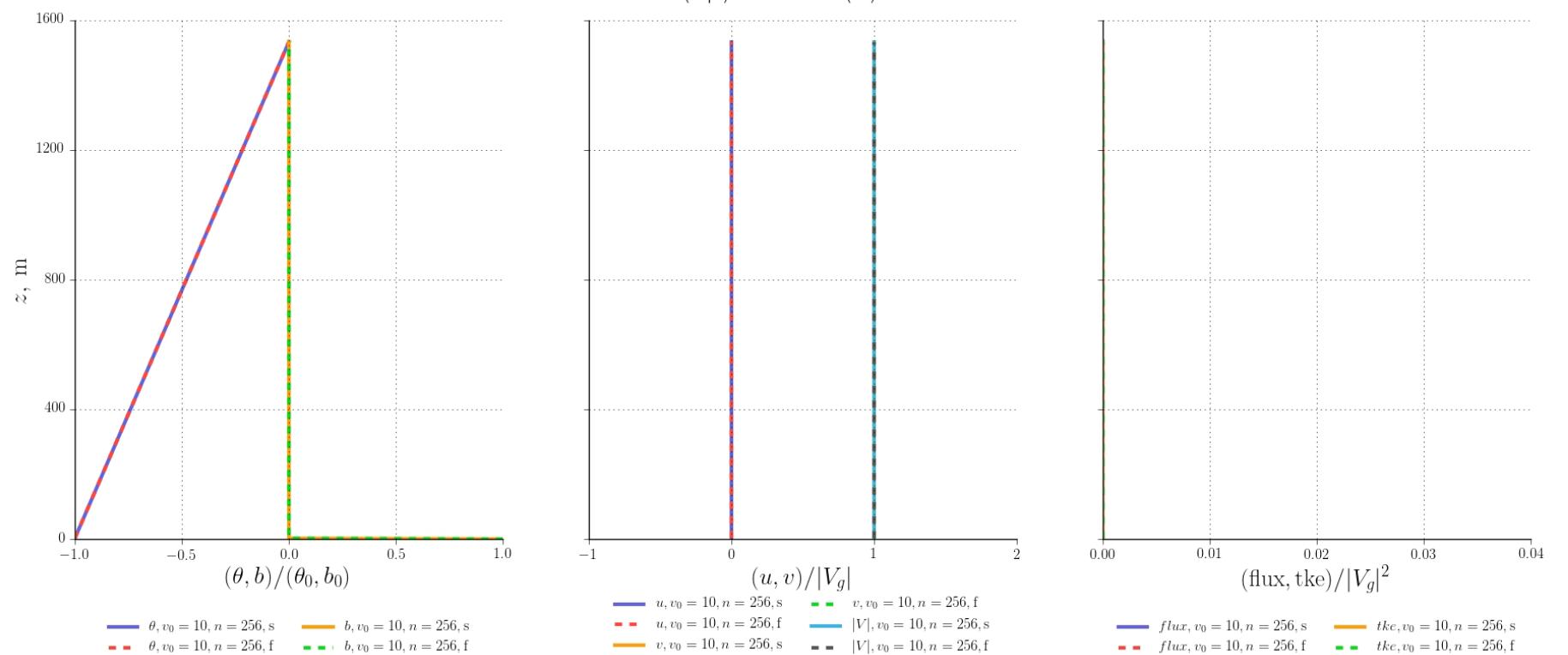
$L_x, L_y = 1024\text{-}4096 \text{ m}$ ;  $N_x, N_y = 256\text{-}1024$ ;  $L_z = 1536 \text{ m}$ ,  $N_z = 384$

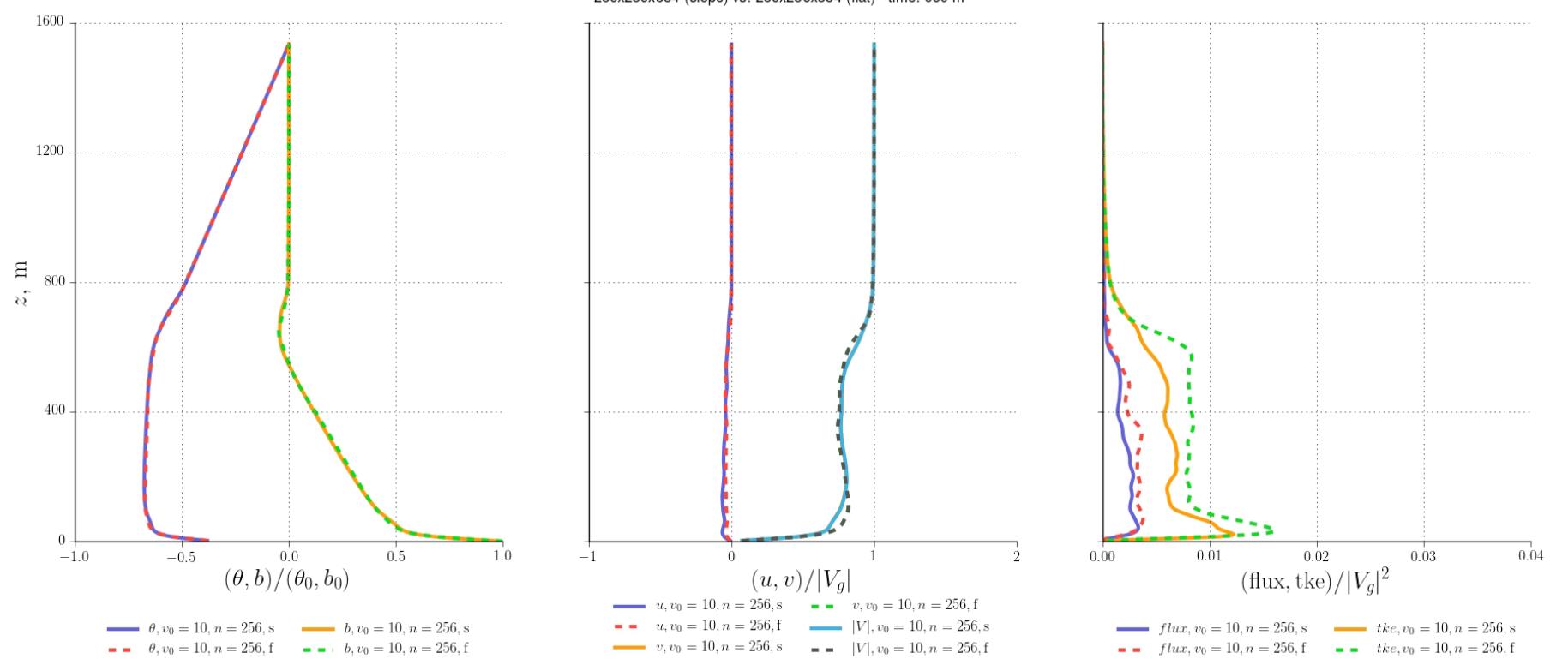
**Numerics:** Micro HH (4<sup>3</sup>), developed by Chiel van Heerwaarden, with 4<sup>th</sup> order Runge-Kutta time integration, 4<sup>th</sup> order advection, and 4<sup>th</sup> order diffusion.

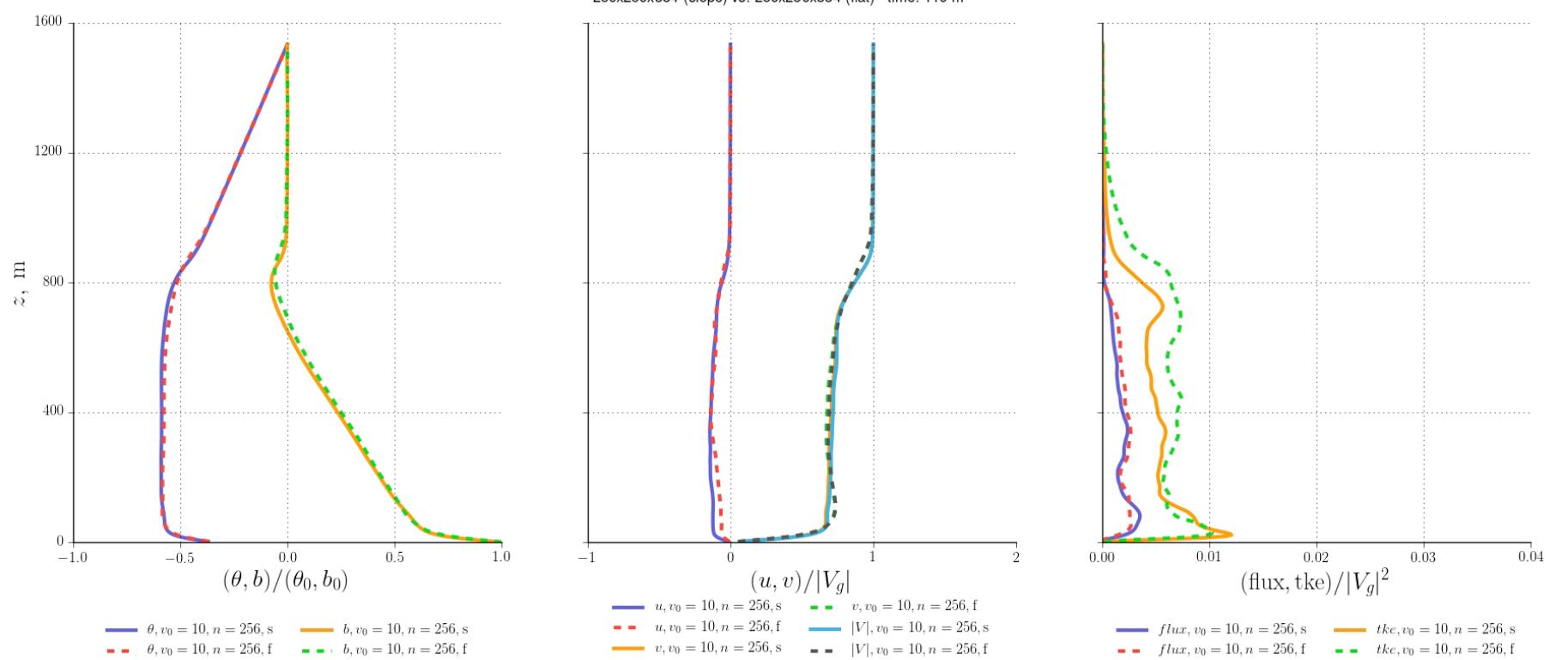
# Slope-angle sensitivity of TKE and $V$ : $\alpha = 0^\circ, 0.09^\circ, 0.18^\circ, 0.27^\circ$

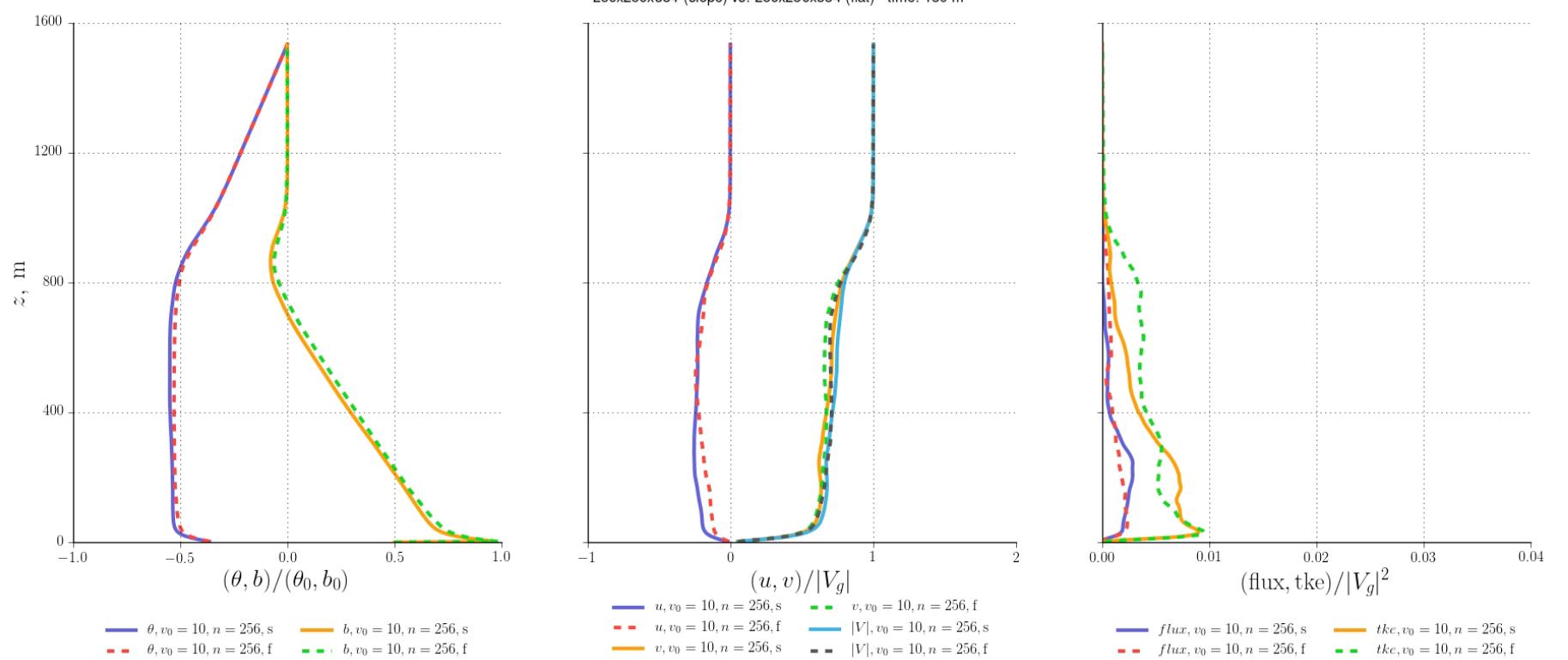
## Surface buoyancy forcing case

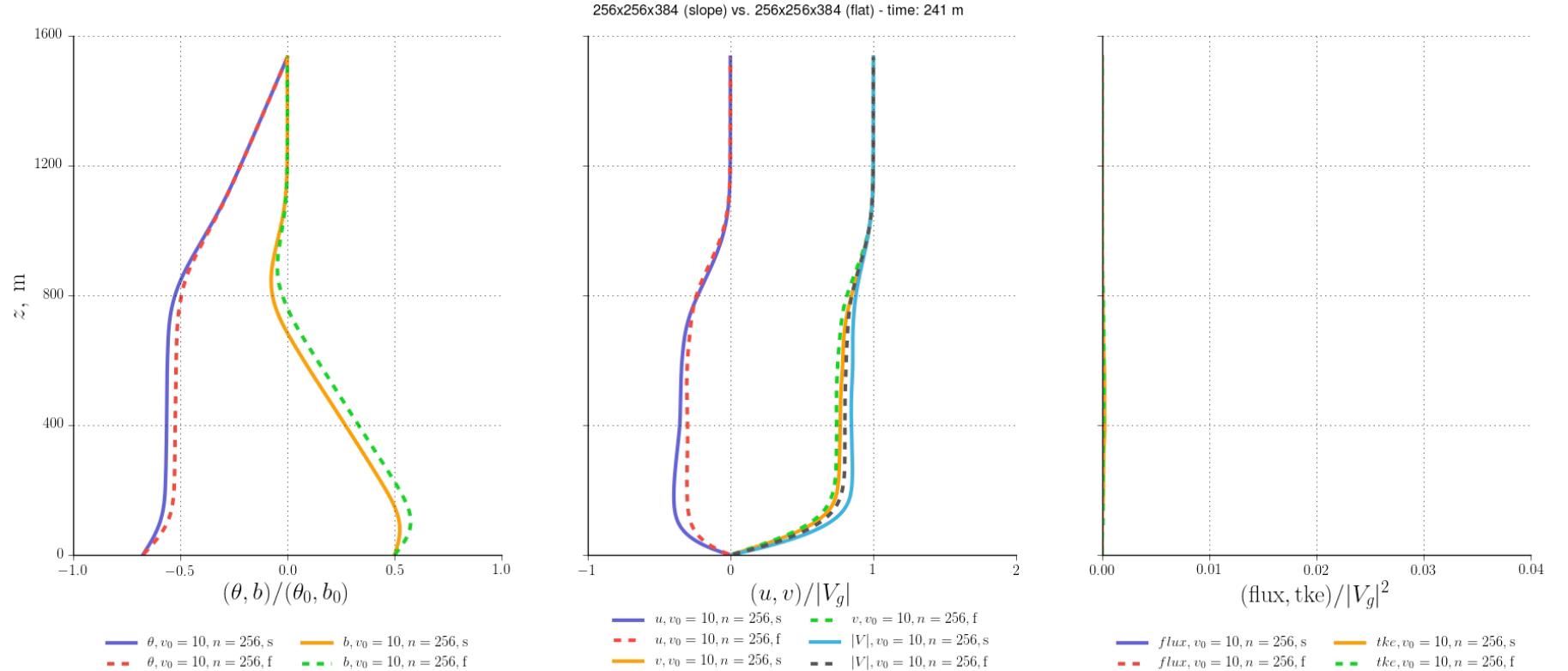


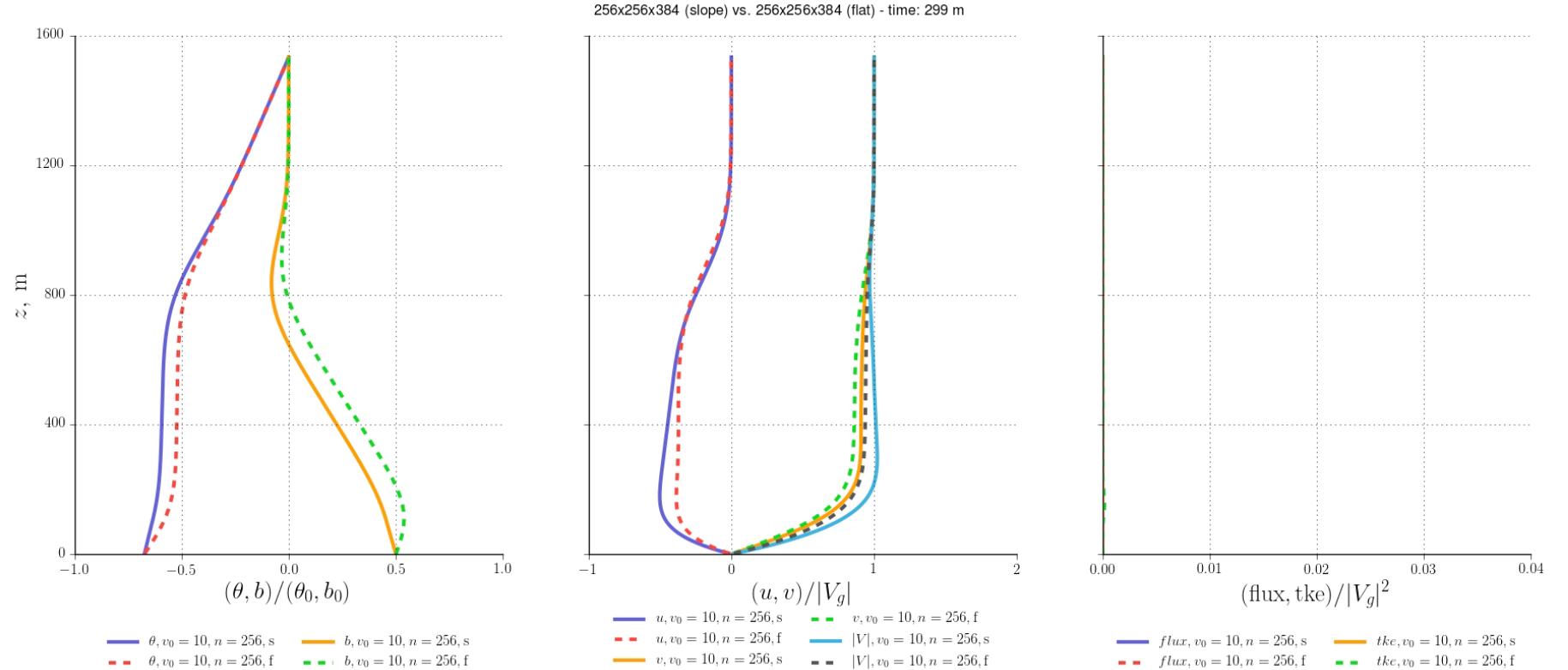


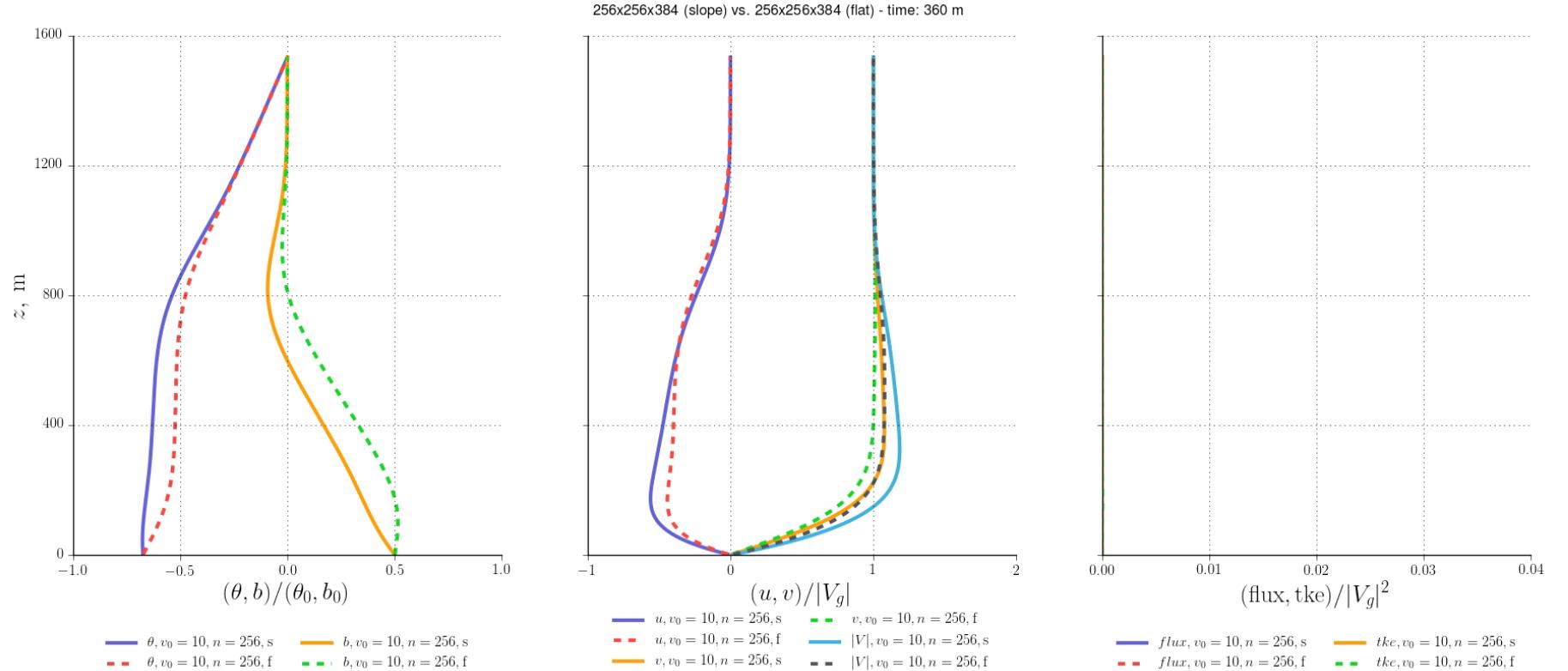


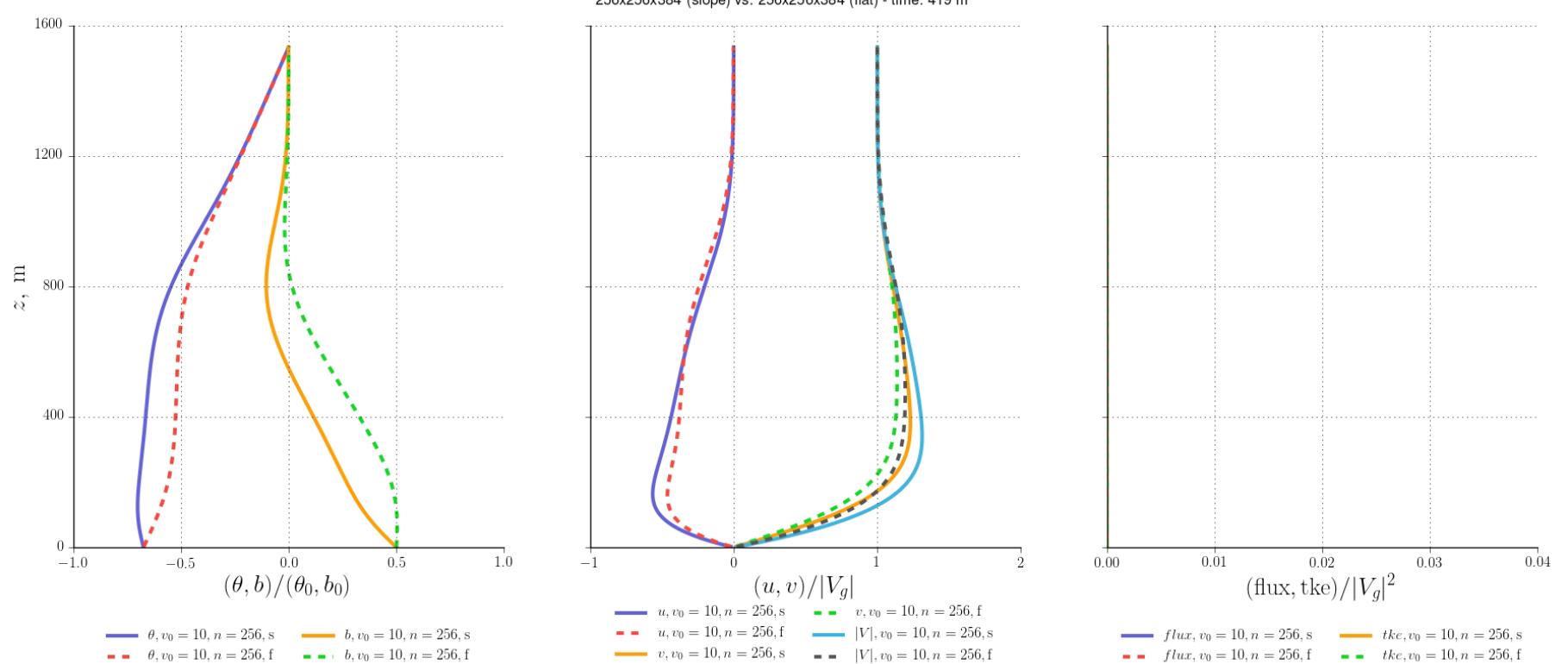


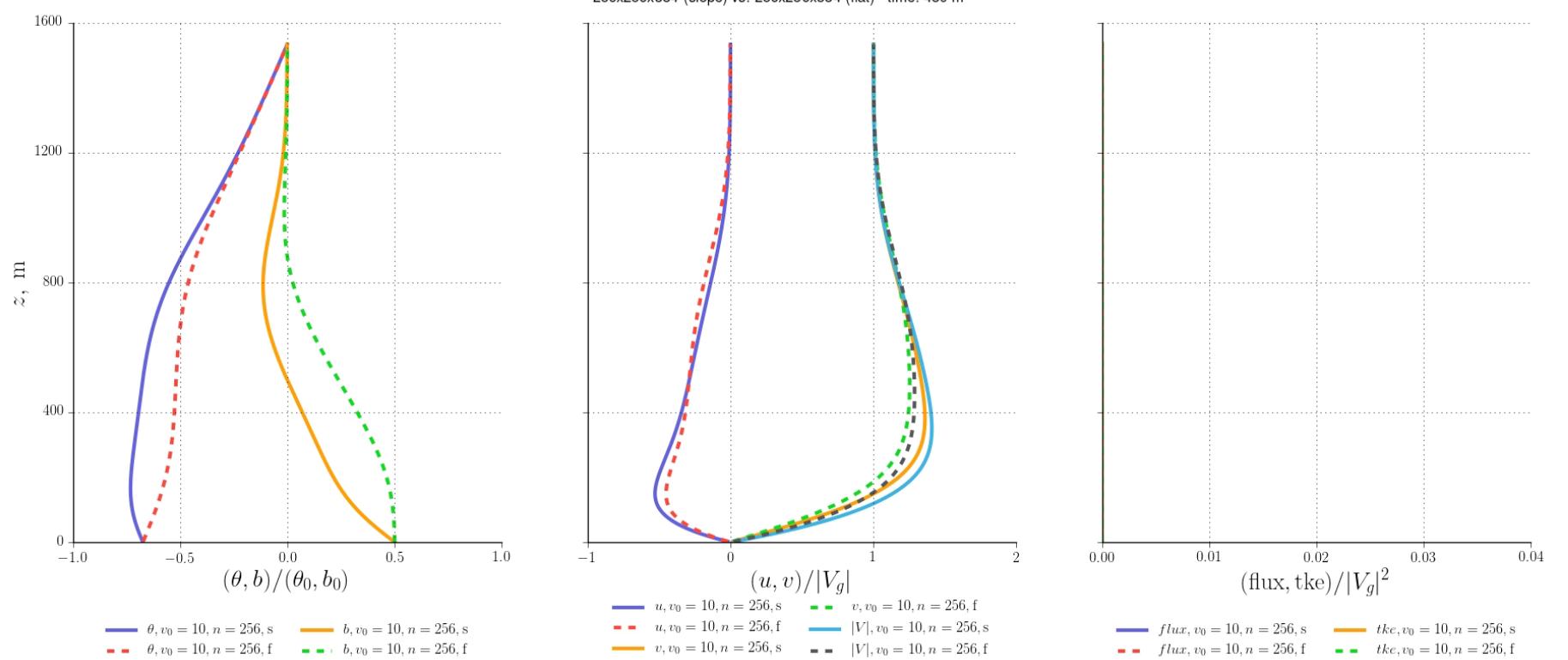


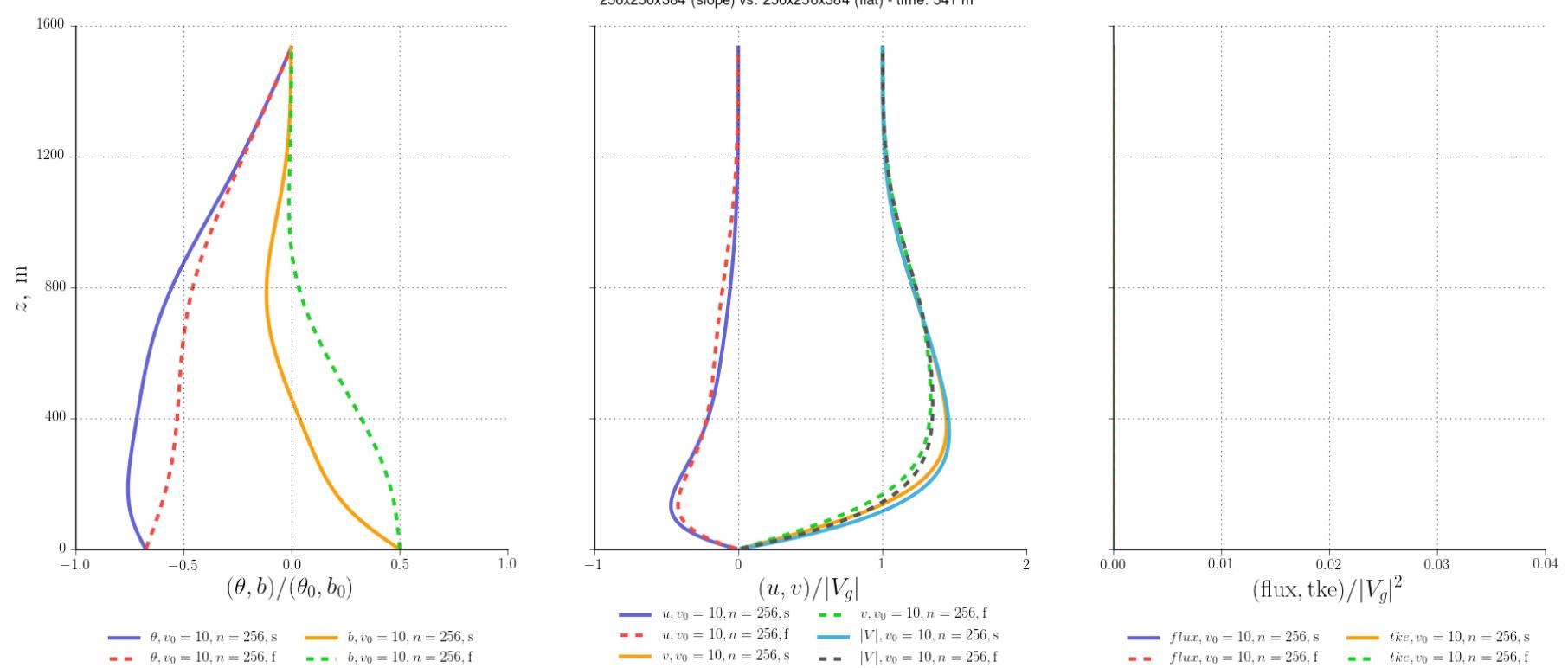


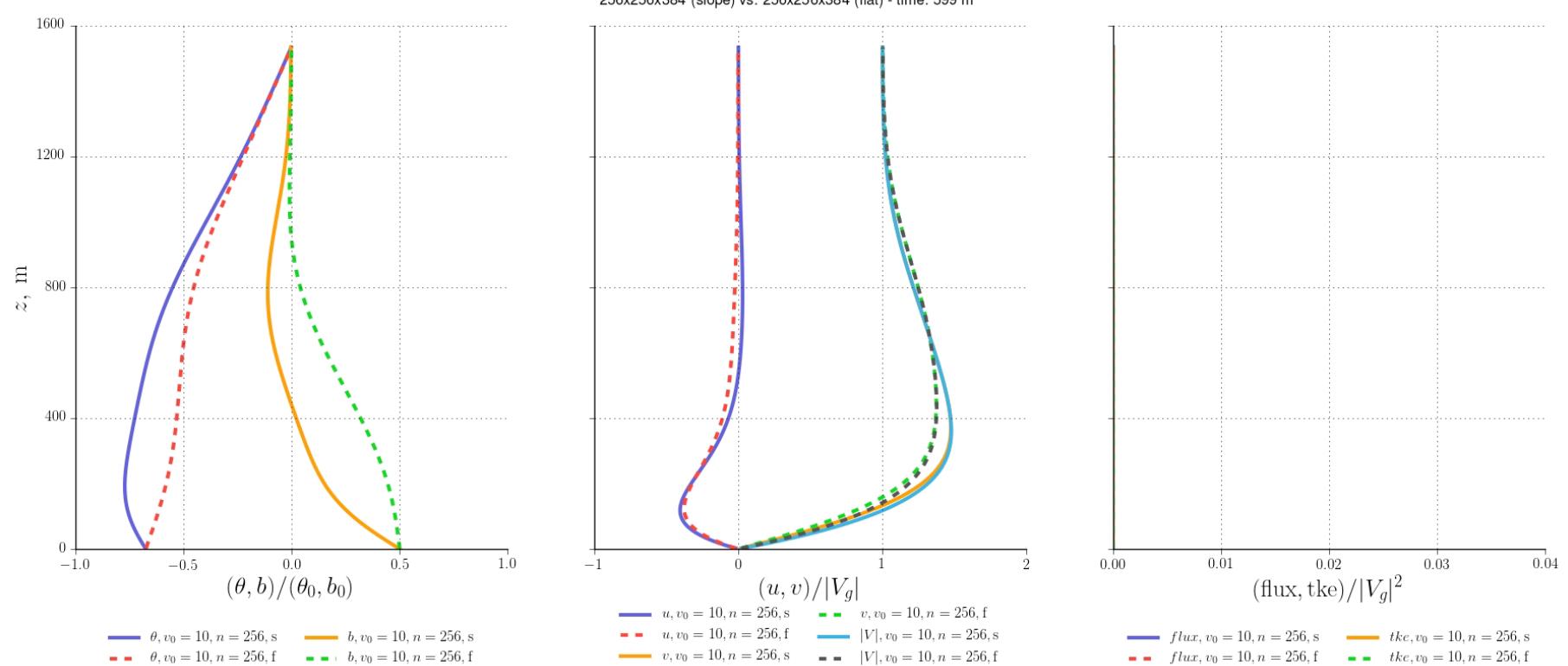


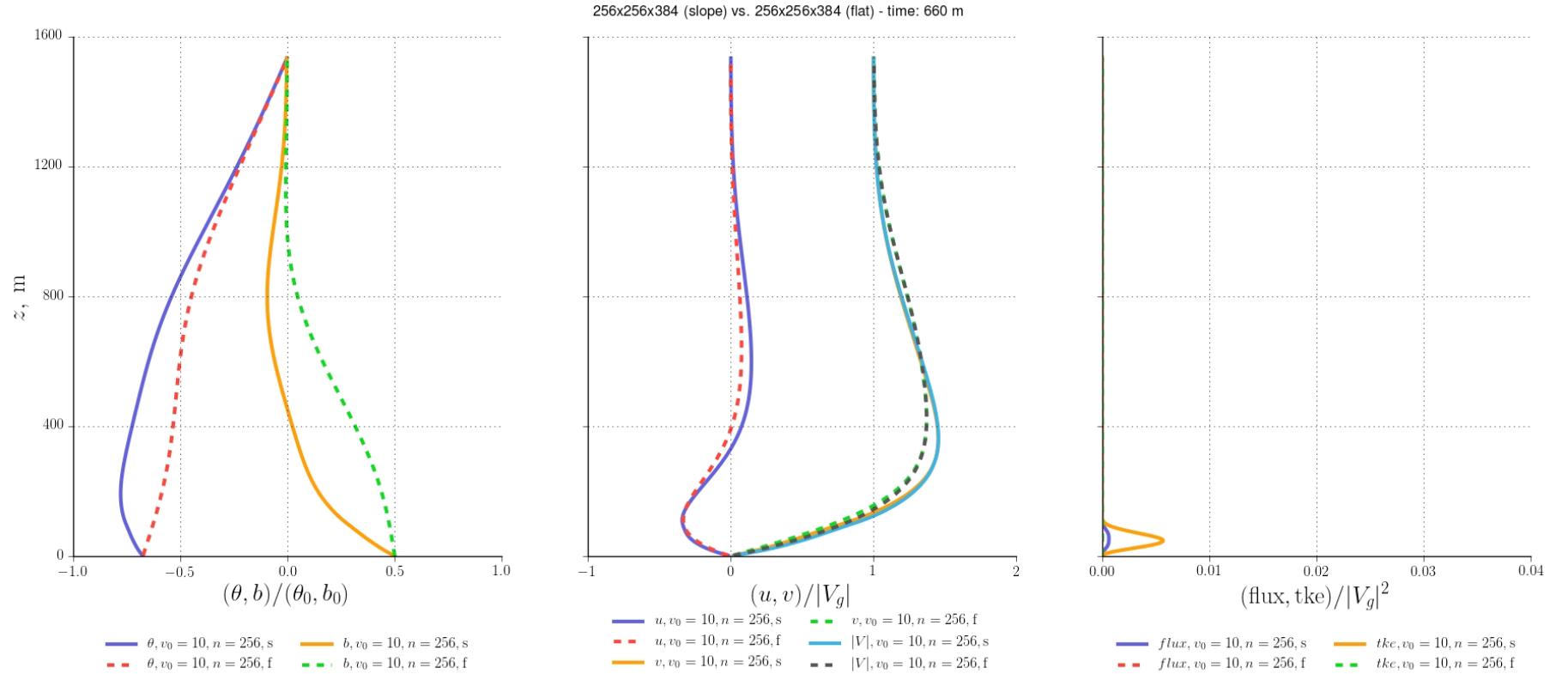


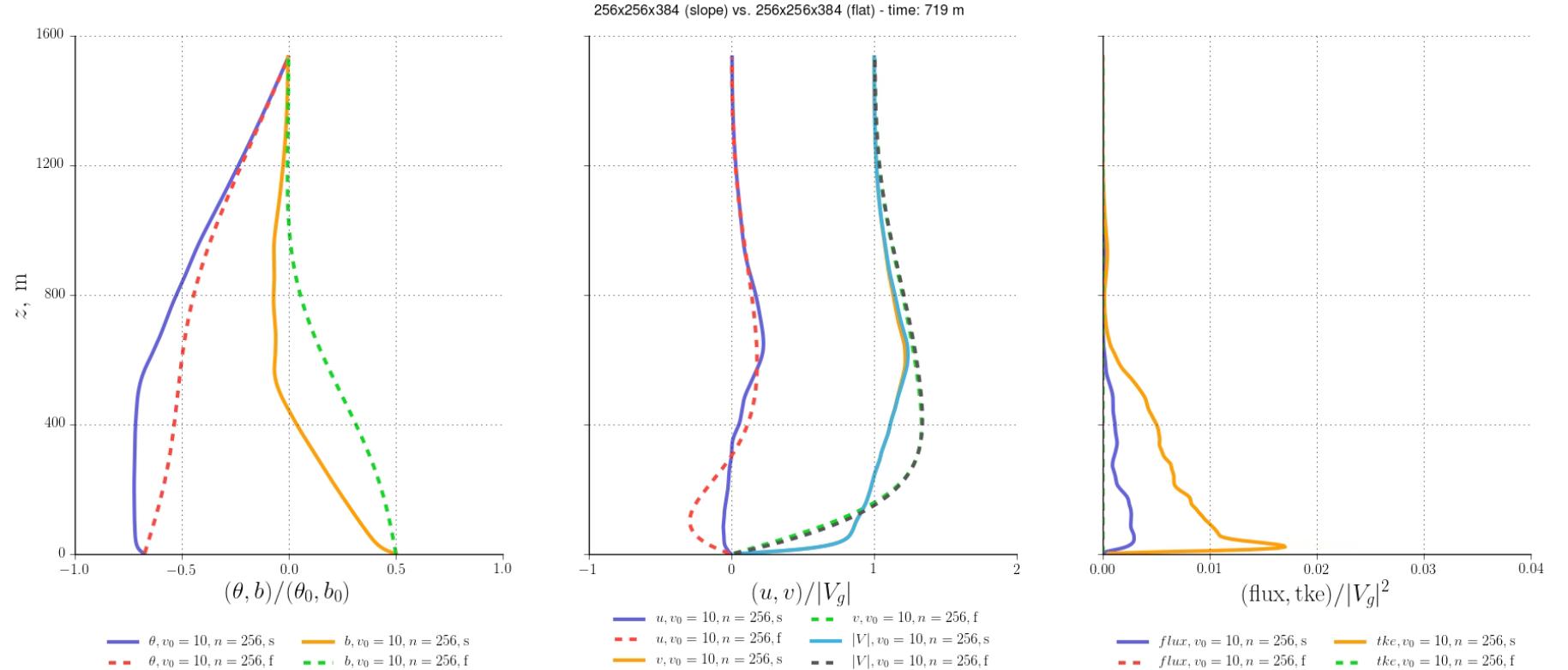


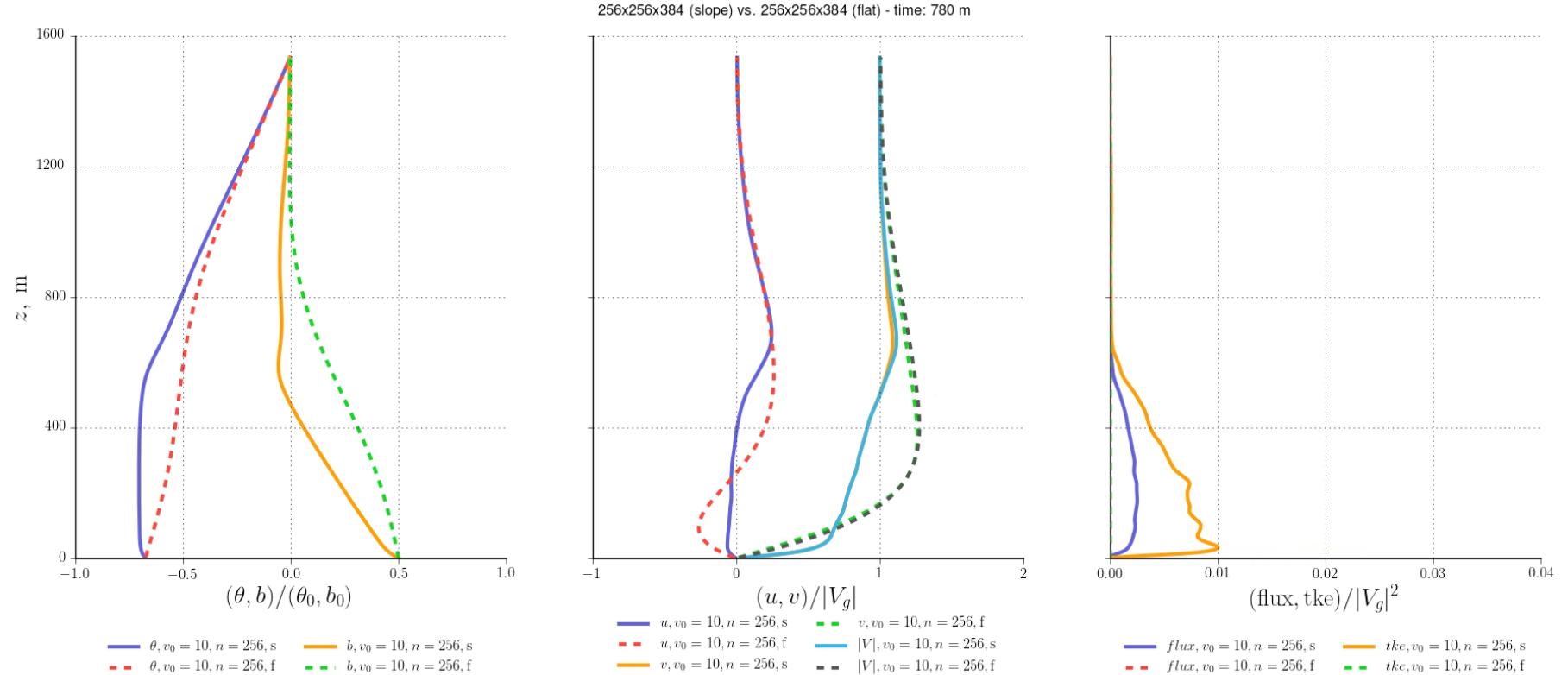


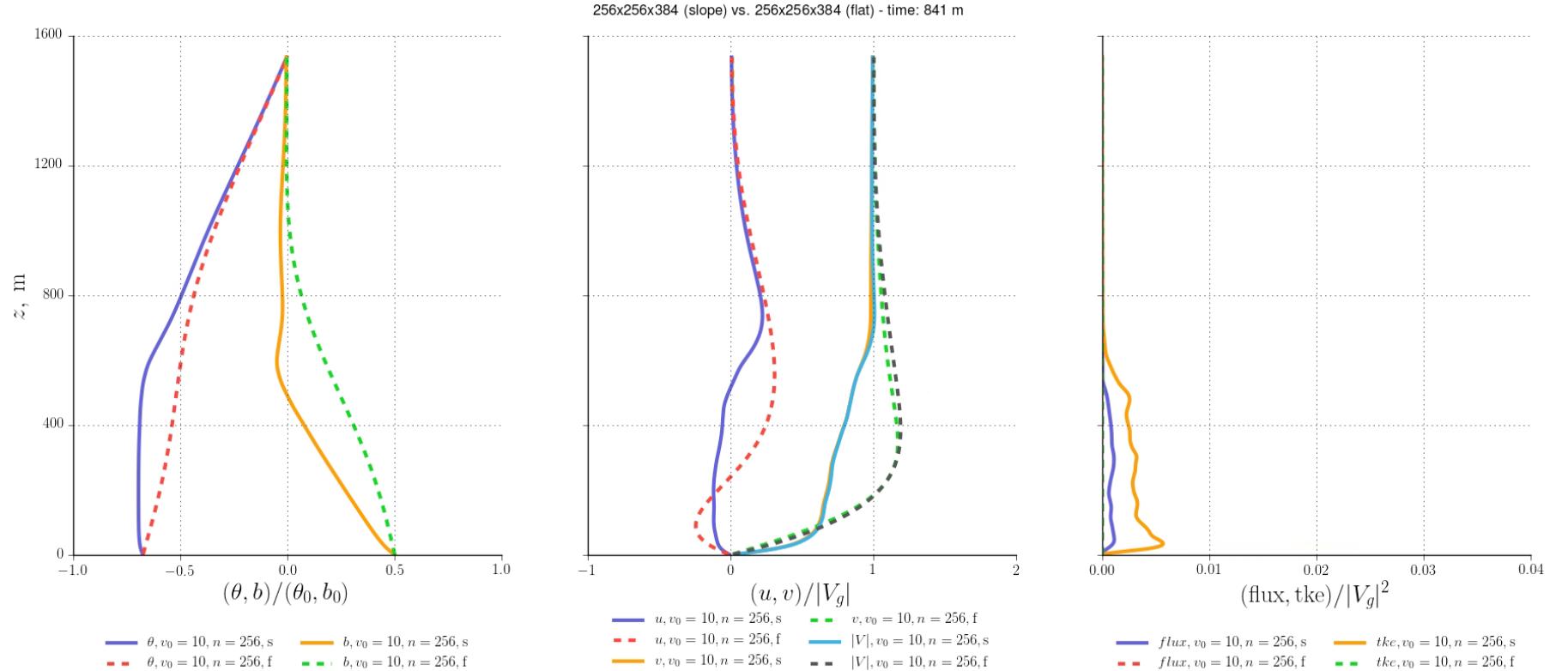


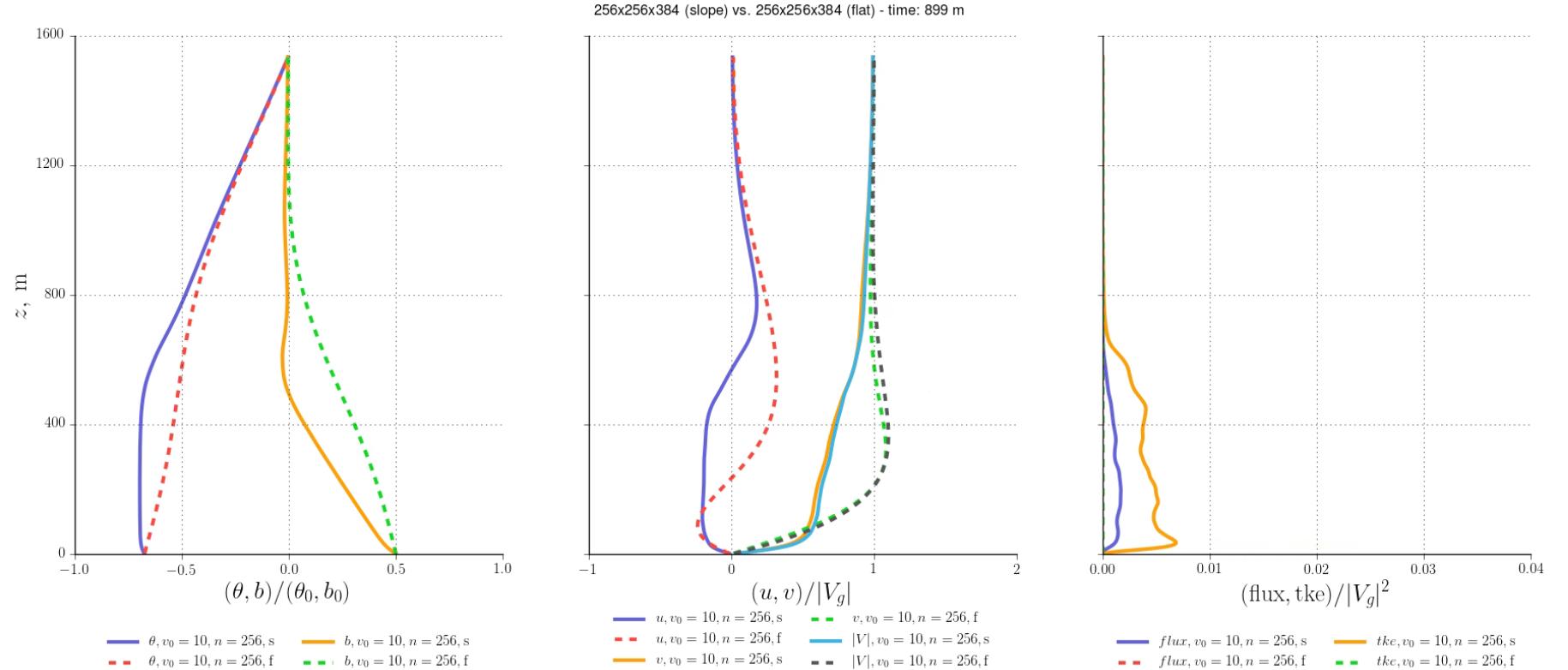






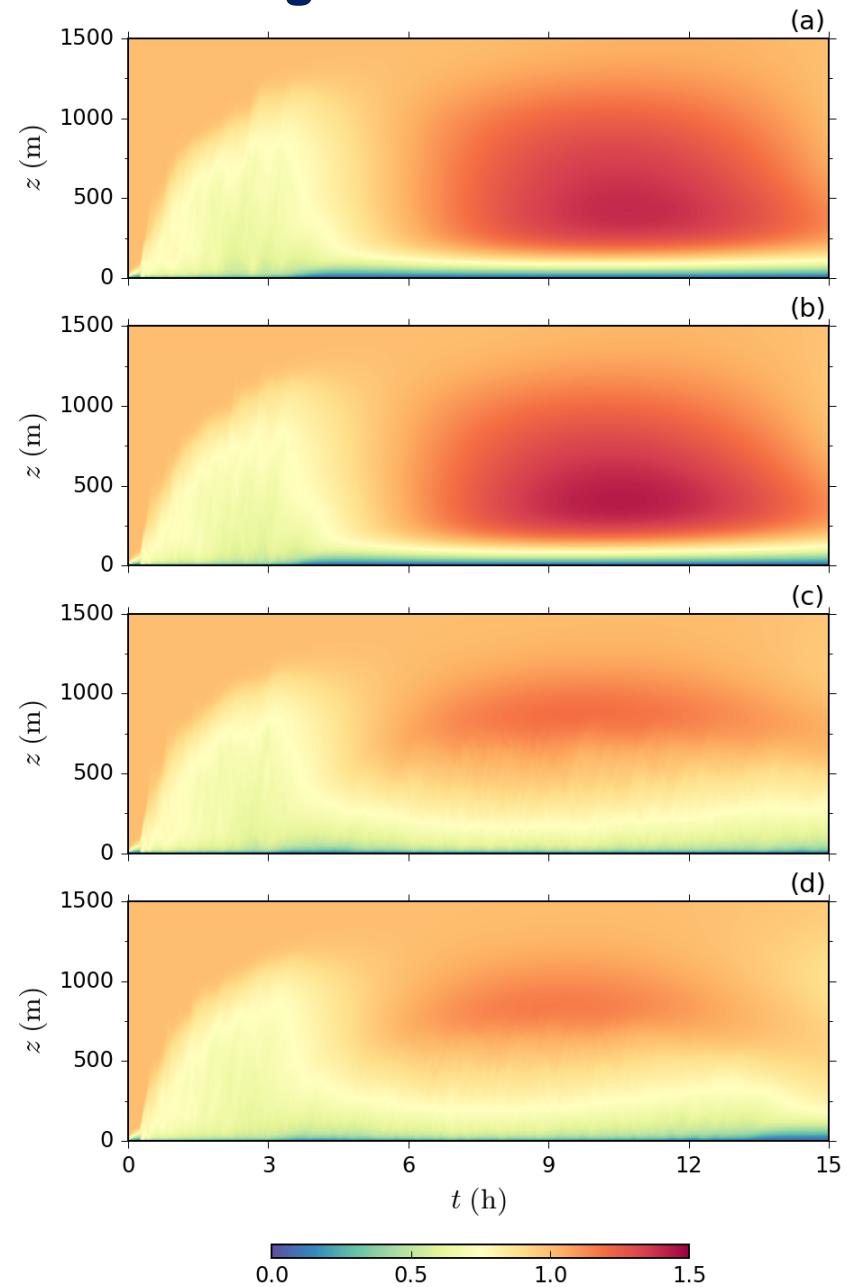
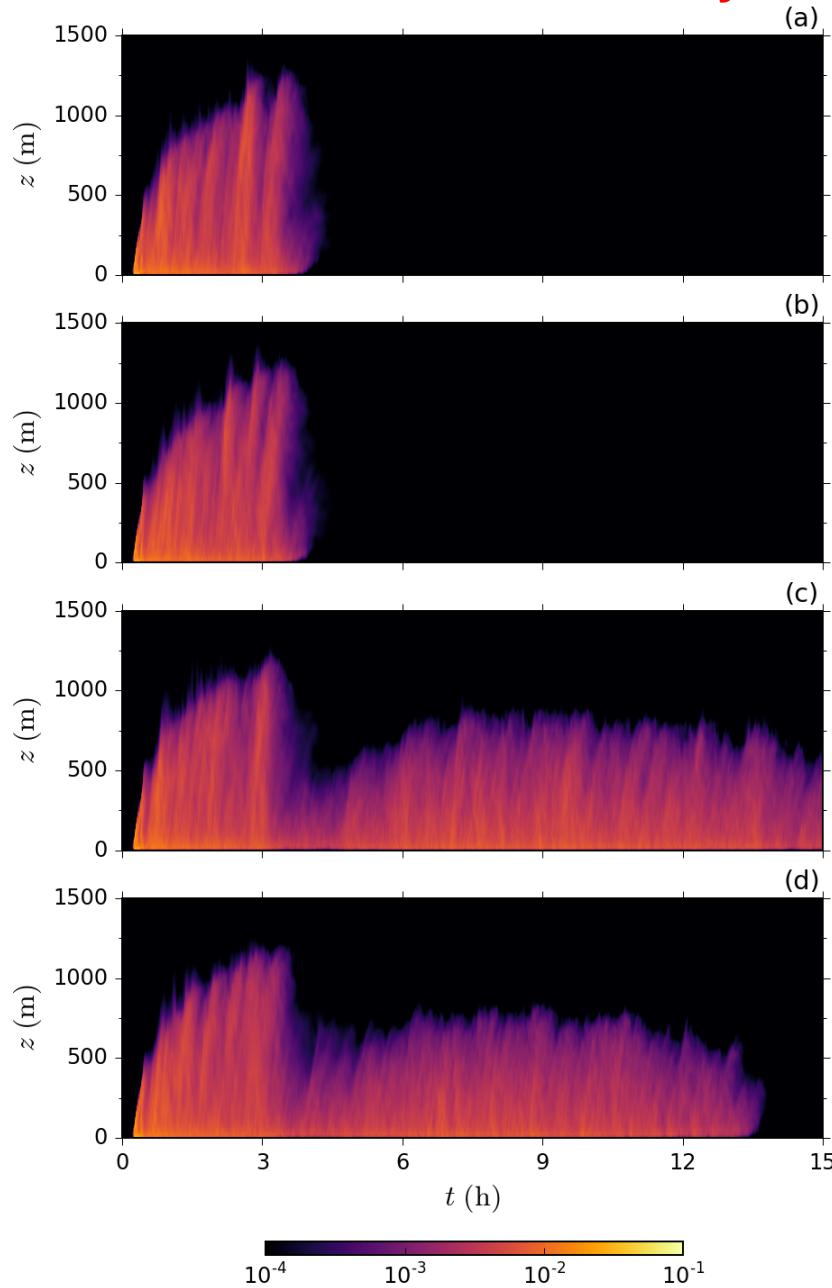


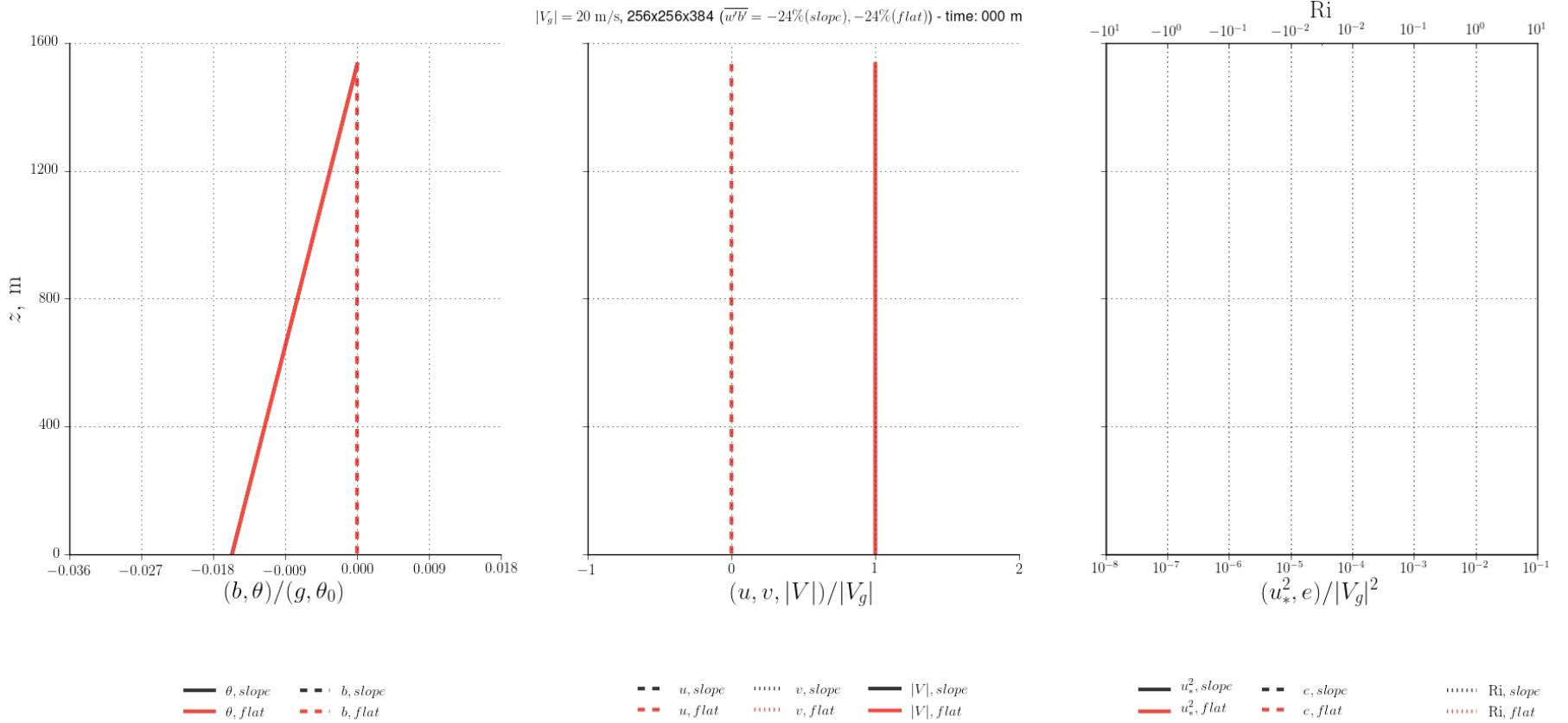


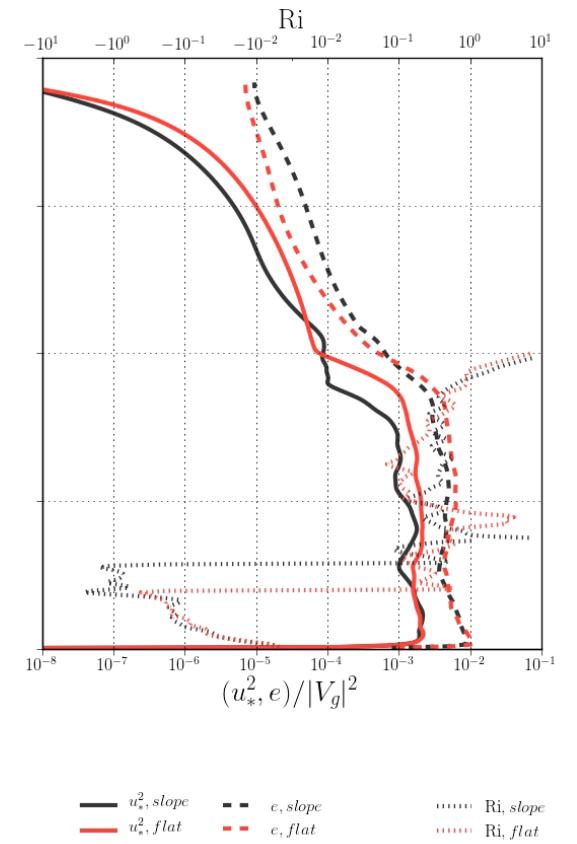
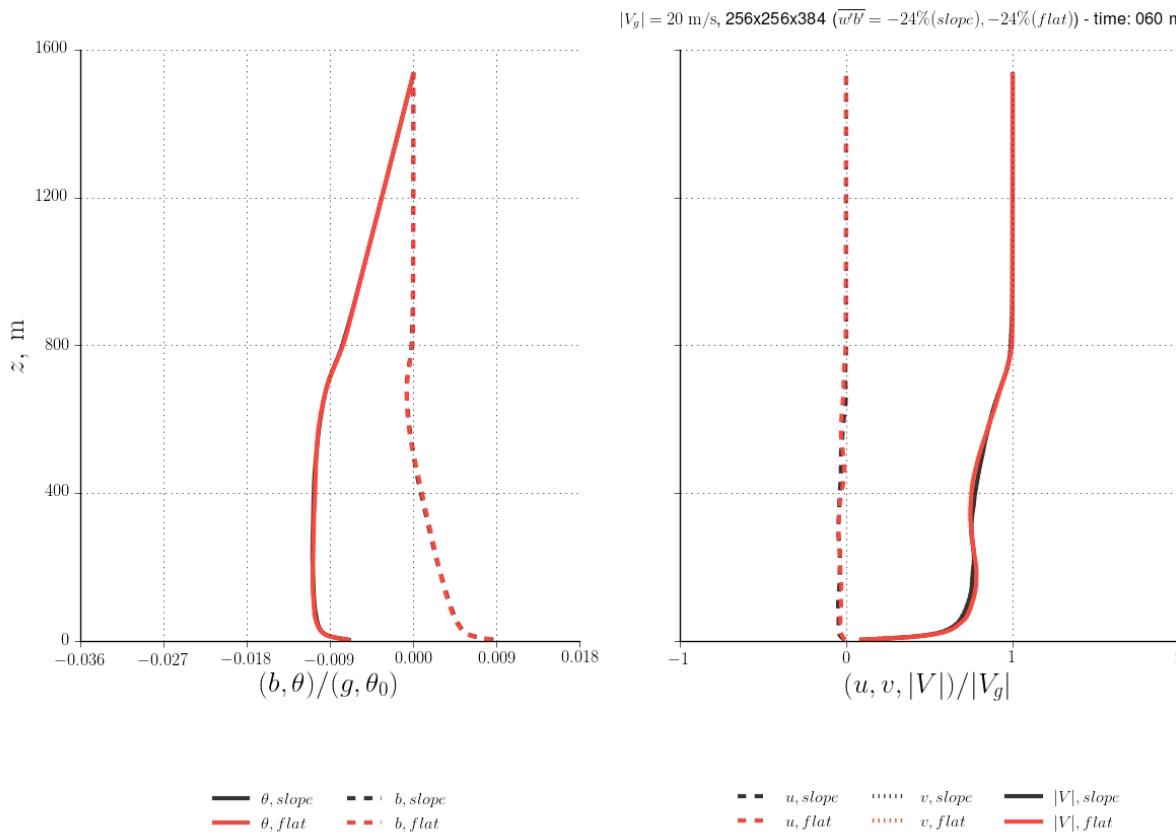


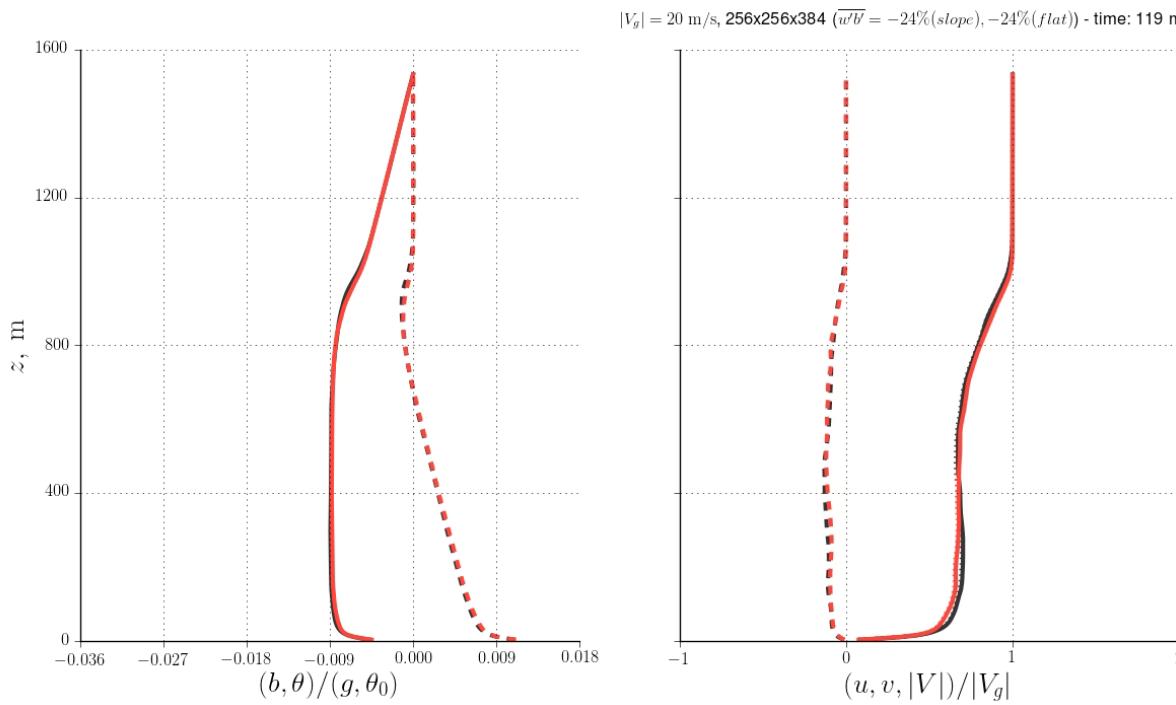
# Slope-angle sensitivity of TKE and $V$ : $\alpha = 0^\circ, 0.09^\circ, 0.18^\circ, 0.27^\circ$

## Surface buoyancy flux forcing case



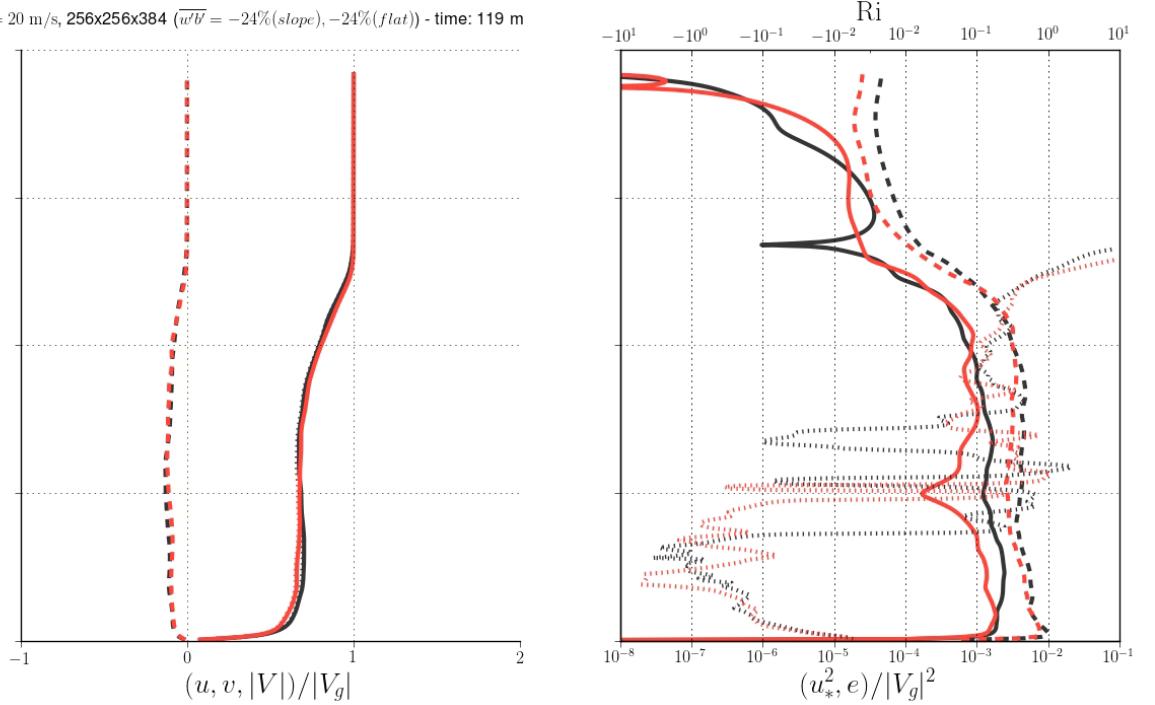






—  $\theta, \text{slope}$   
—  $\theta, \text{flat}$

—  $b, \text{slope}$   
—  $b, \text{flat}$



—  $u, \text{slope}$   
—  $u, \text{flat}$

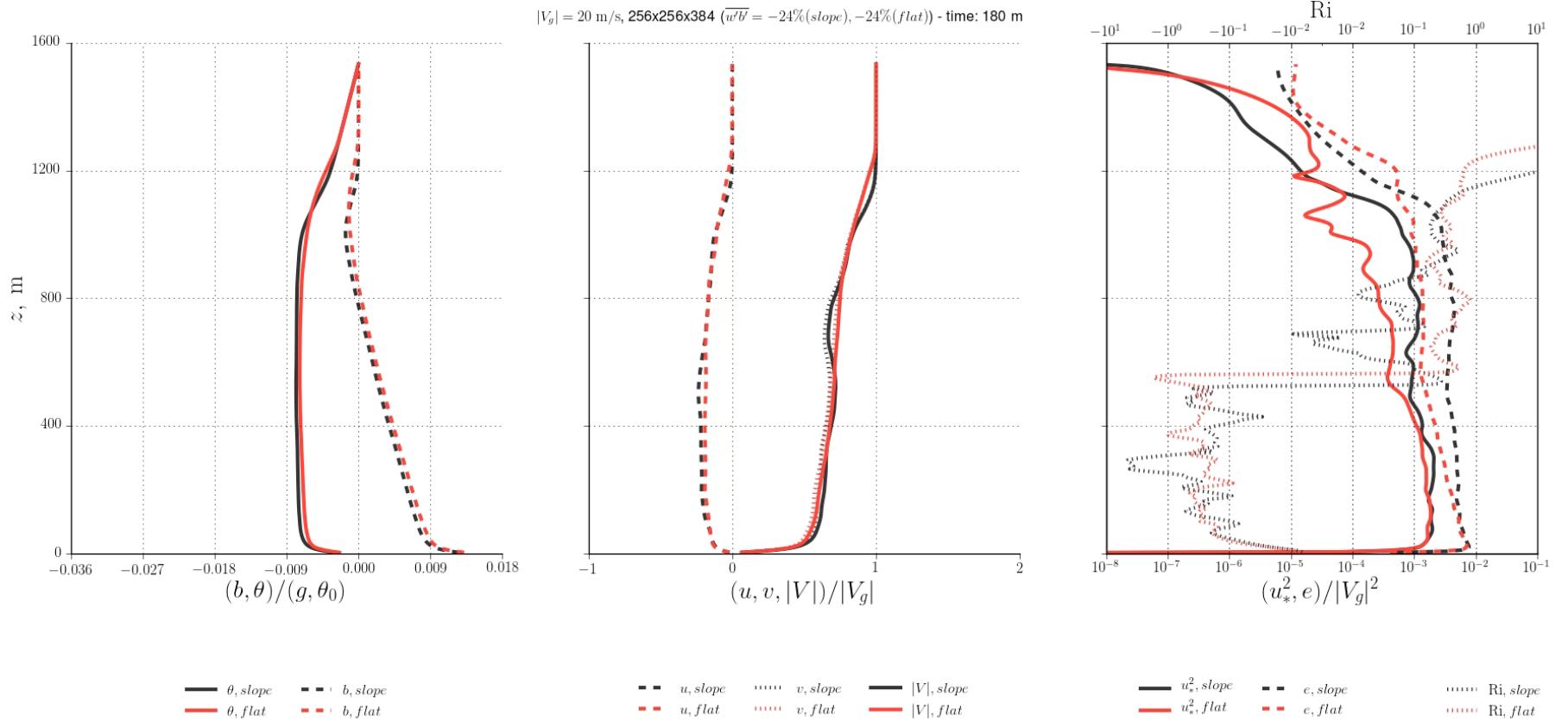
—  $v, \text{slope}$   
—  $v, \text{flat}$

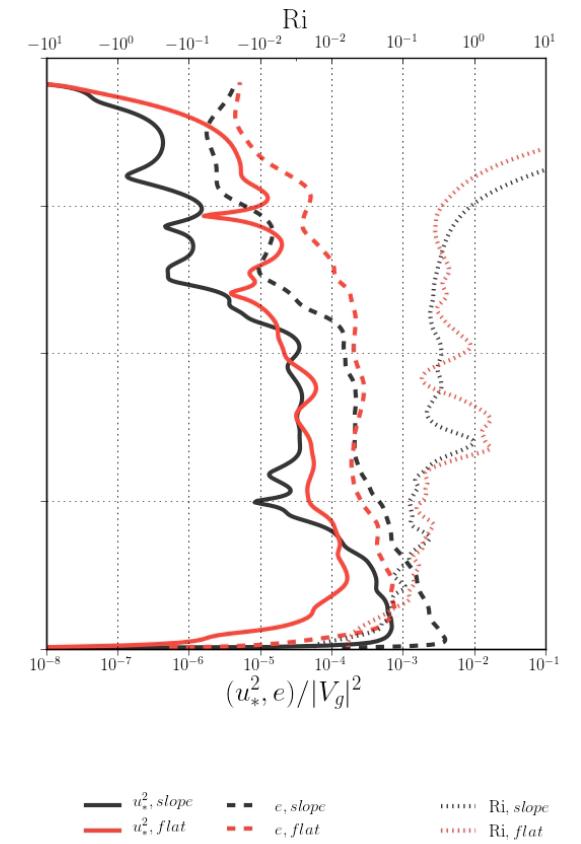
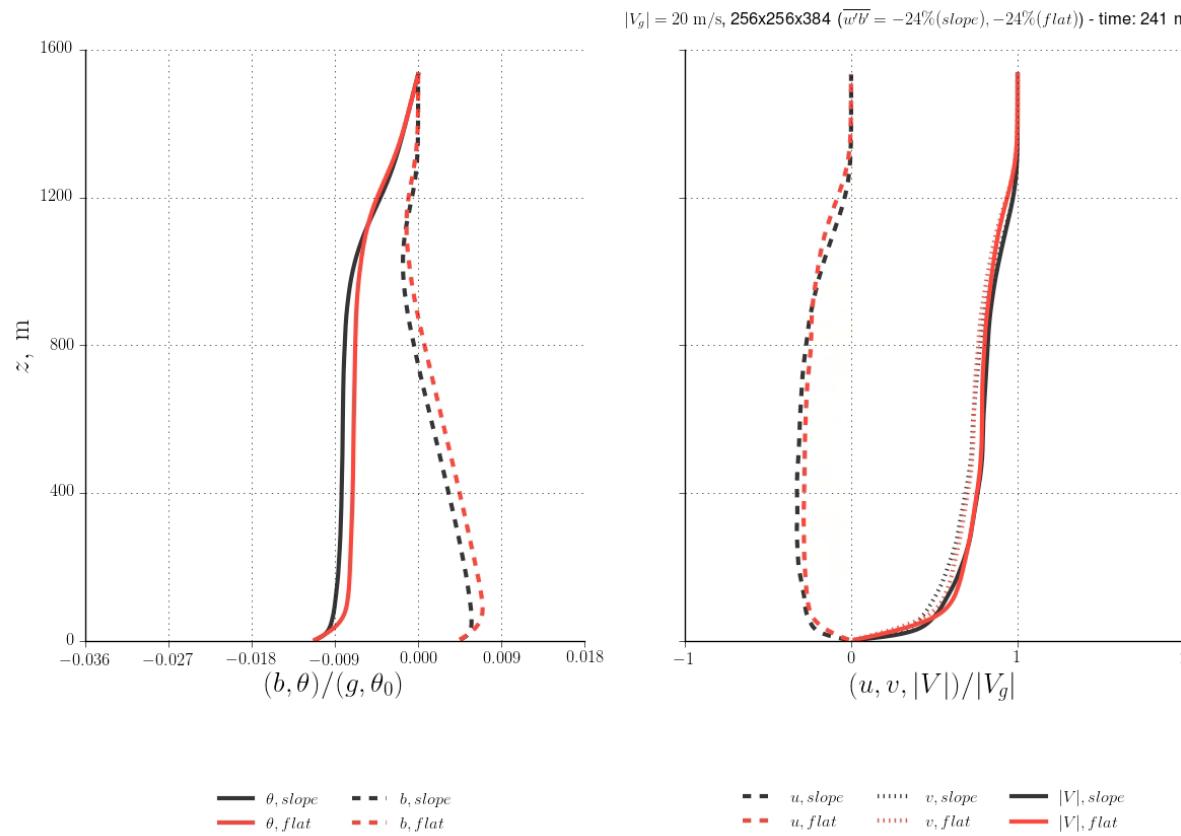
—  $|V|, \text{slope}$   
—  $|V|, \text{flat}$

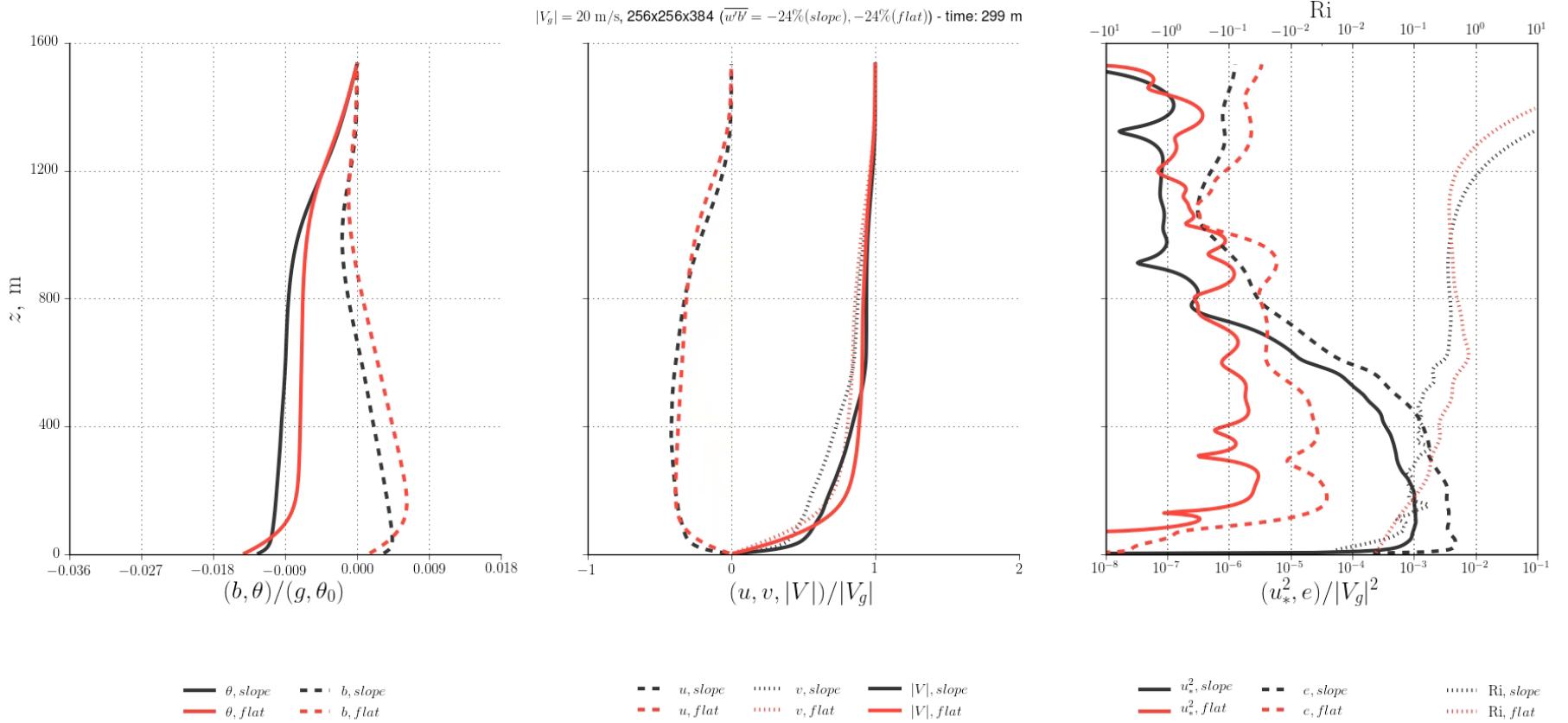
—  $u_*^2, \text{slope}$   
—  $u_*^2, \text{flat}$

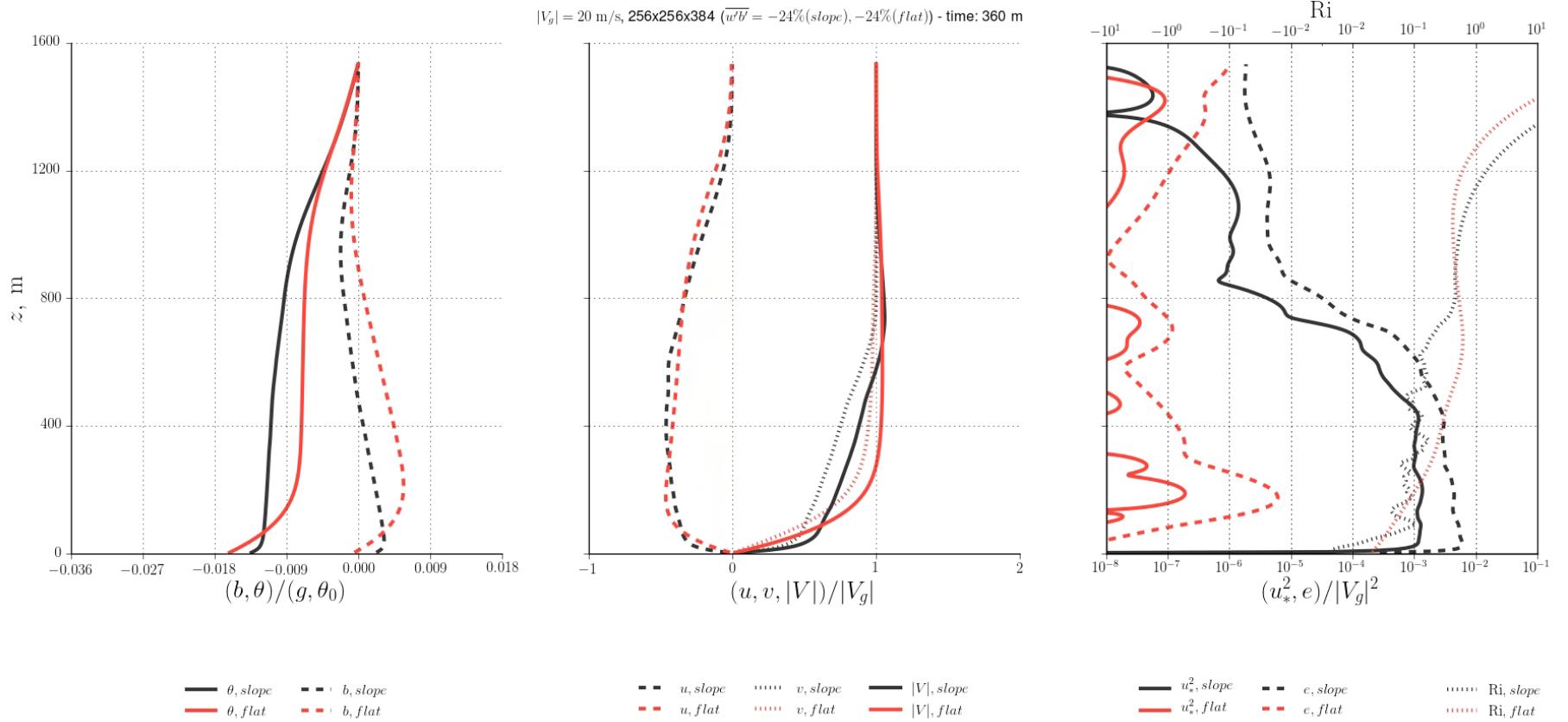
—  $e, \text{slope}$   
—  $e, \text{flat}$

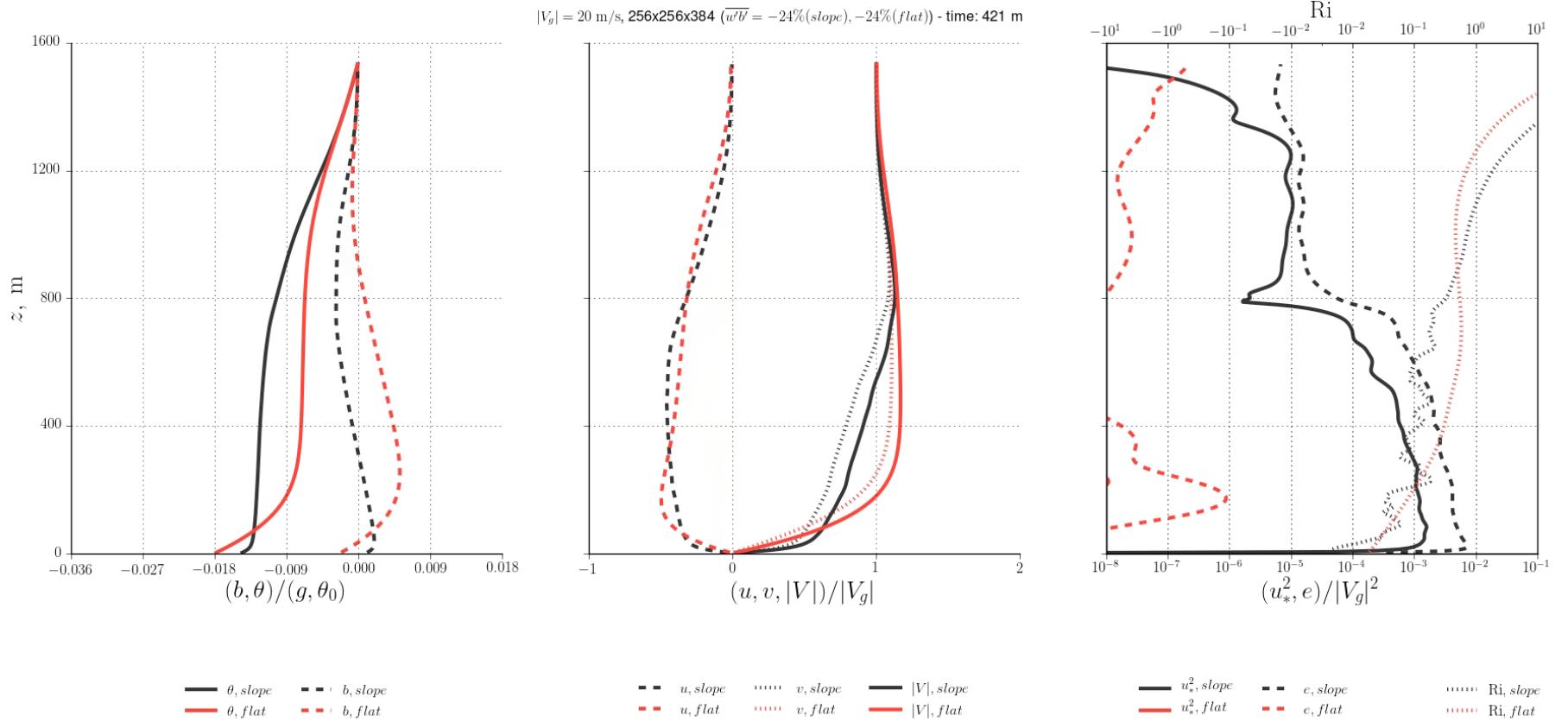
—  $\text{Ri, slope}$   
—  $\text{Ri, flat}$

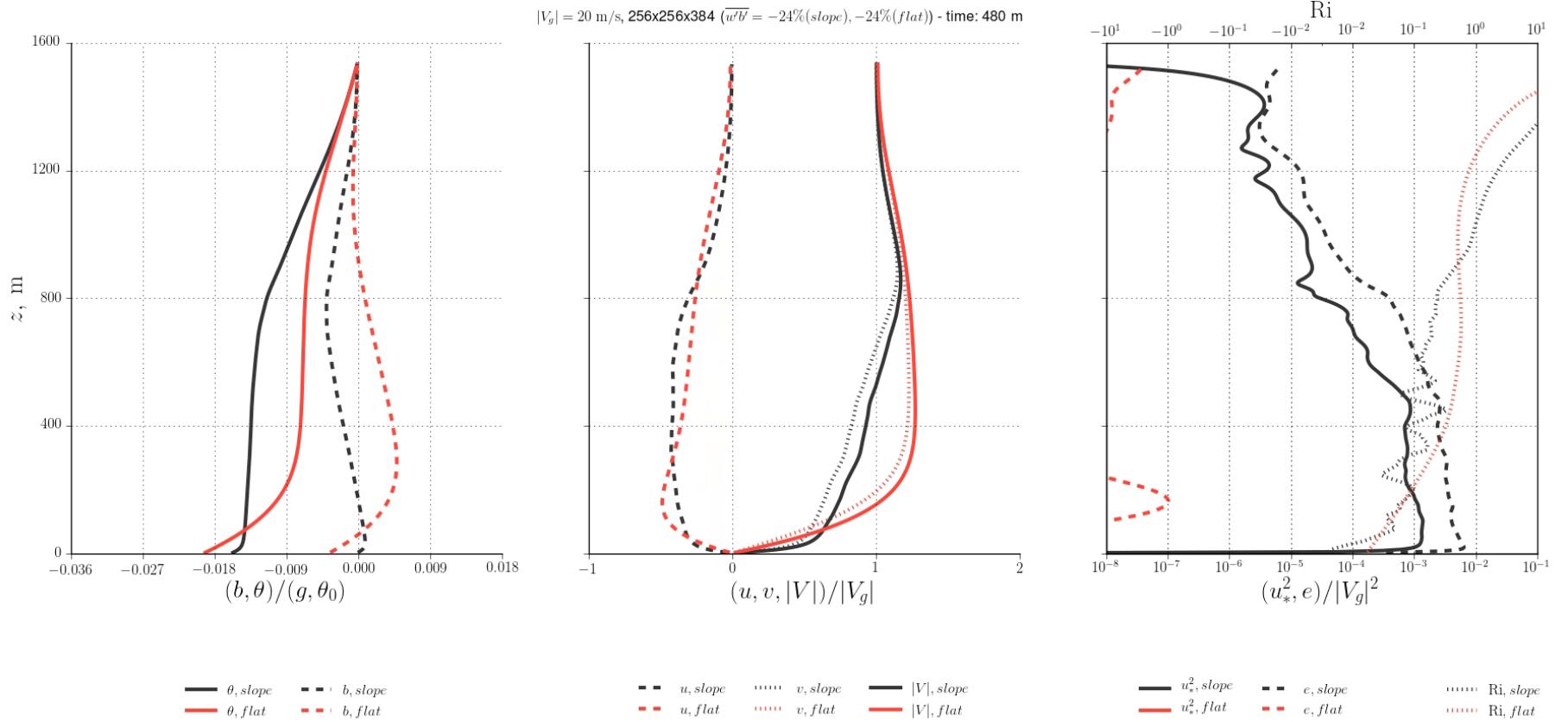


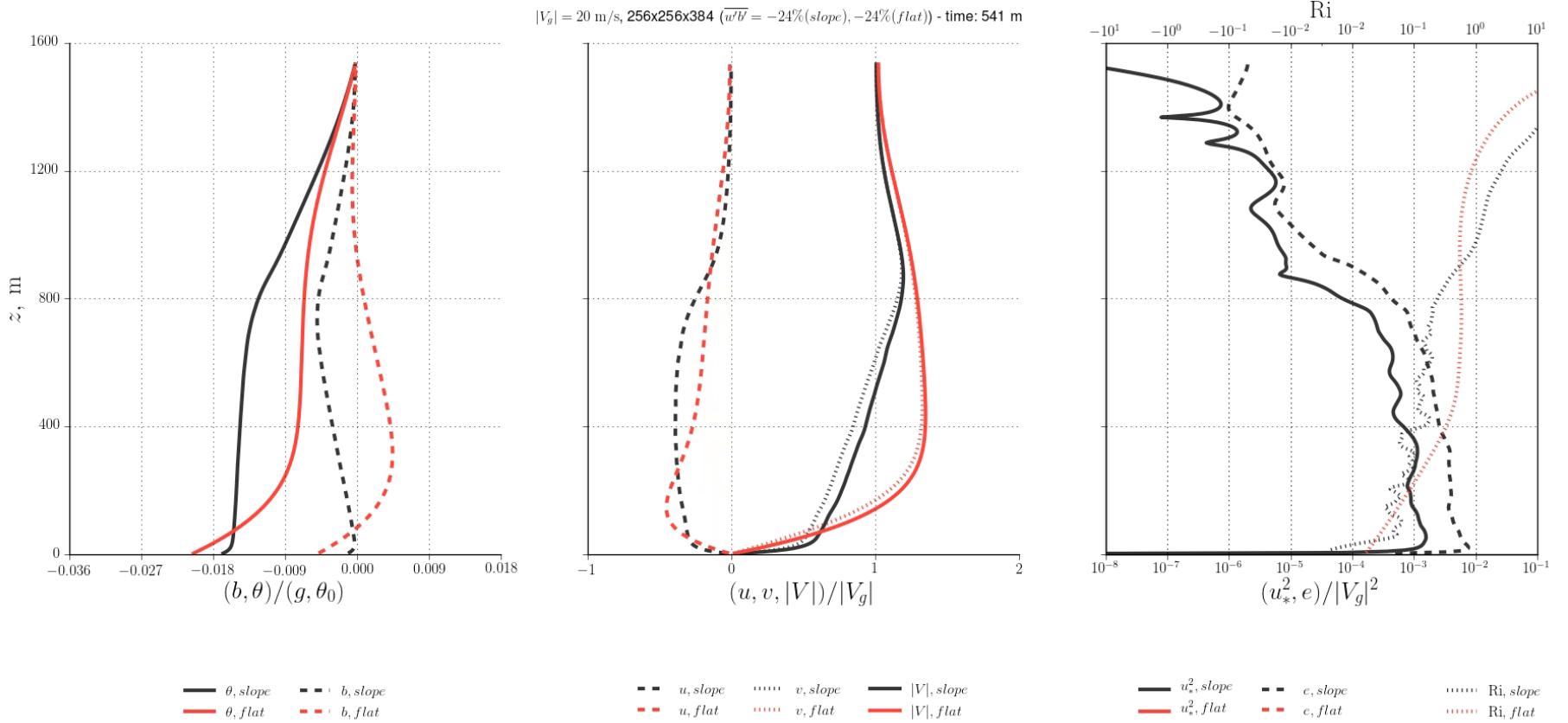


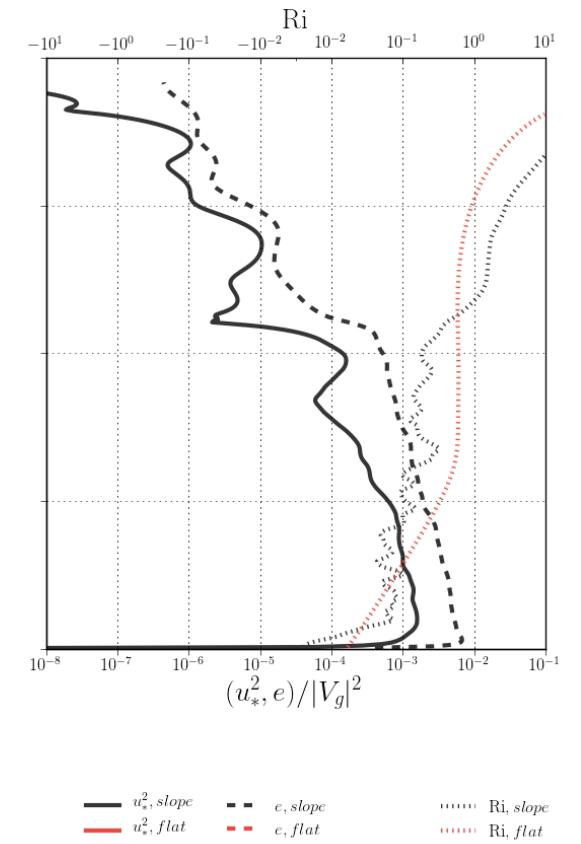
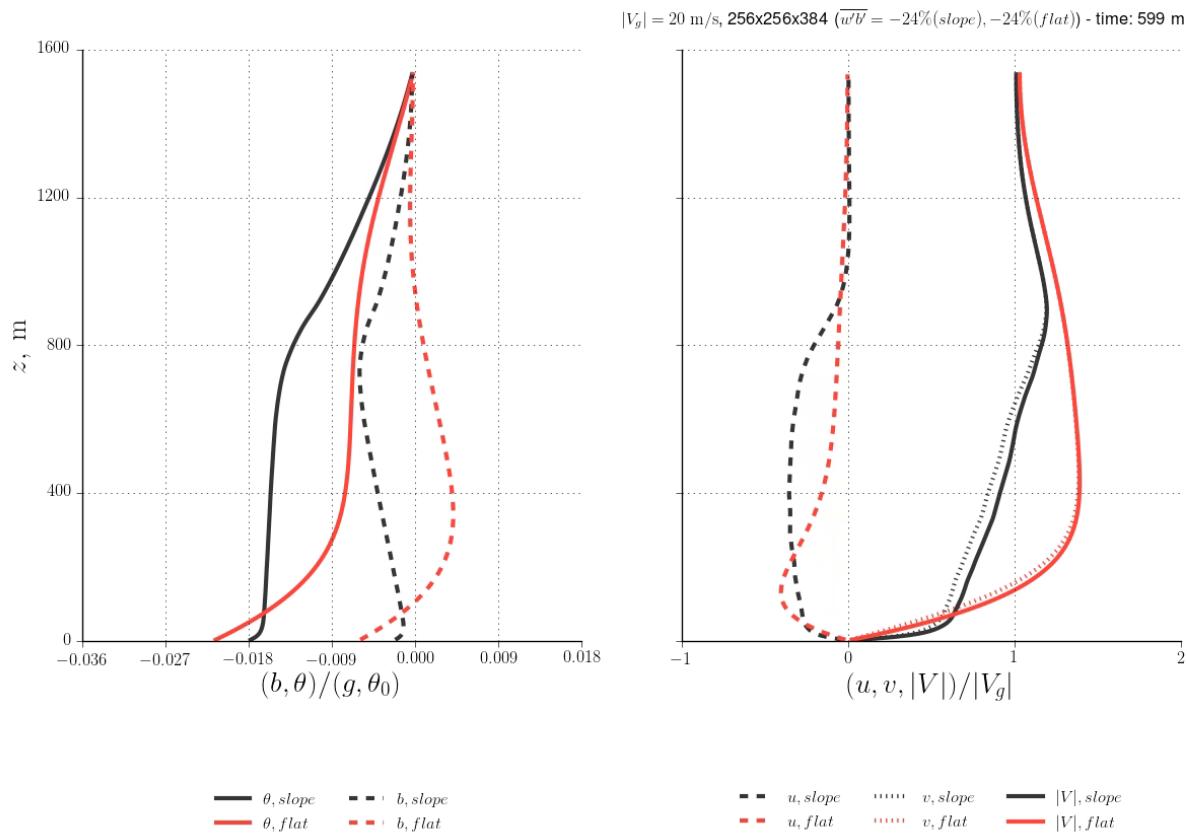


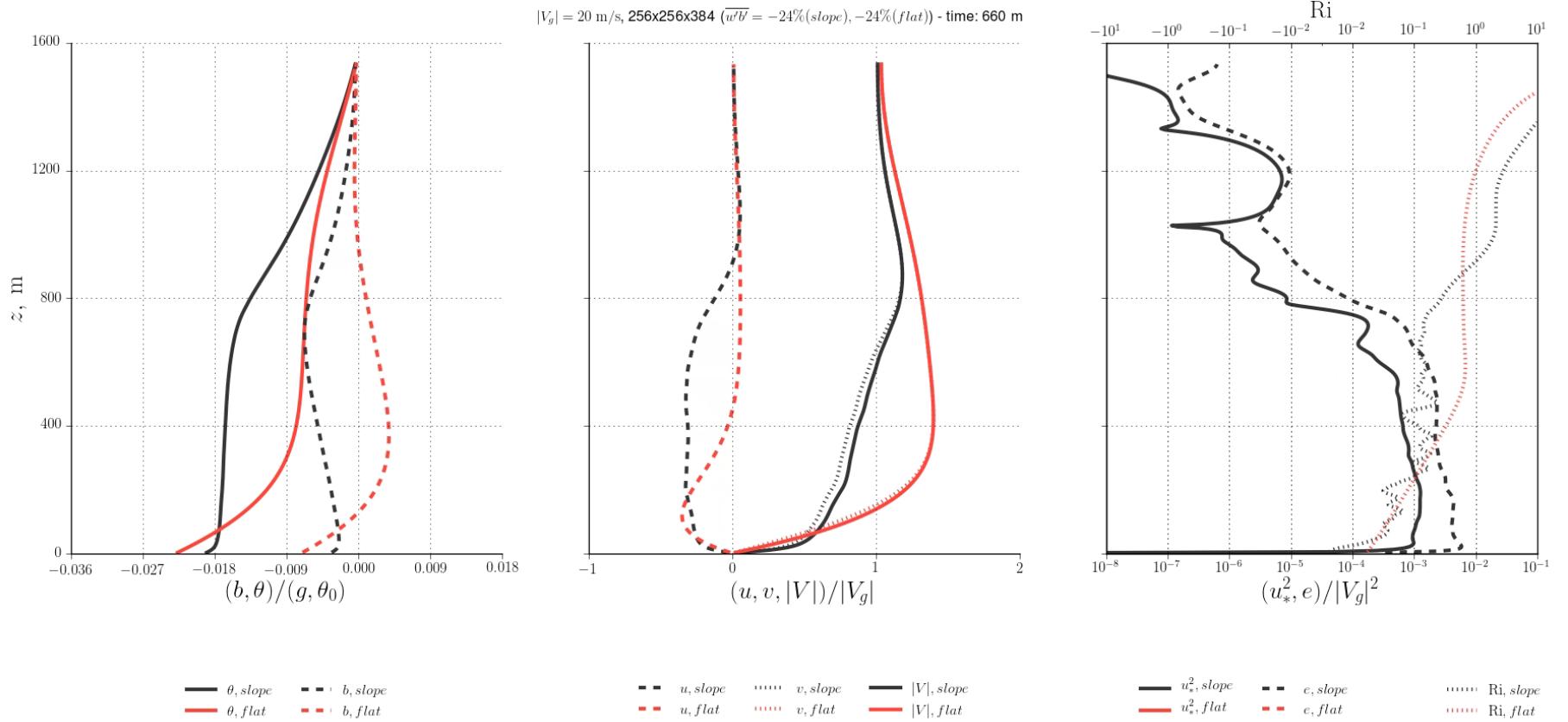


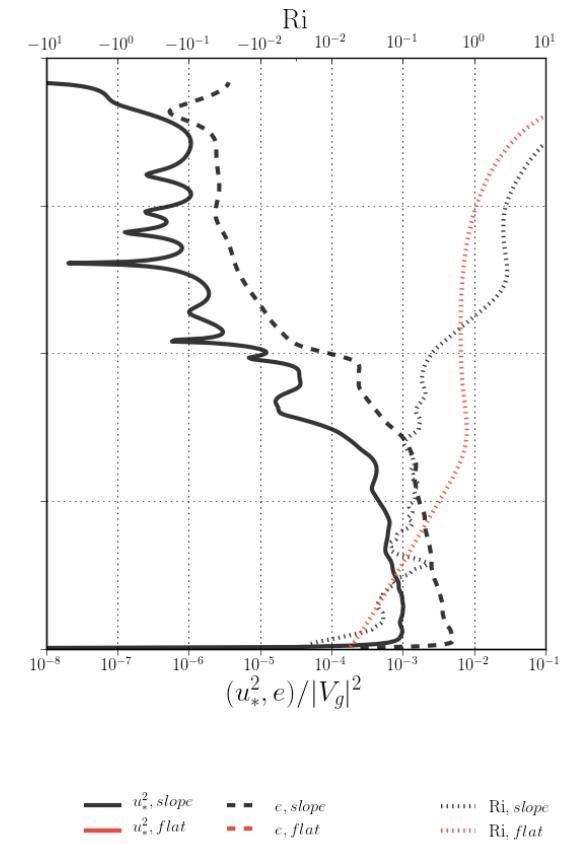
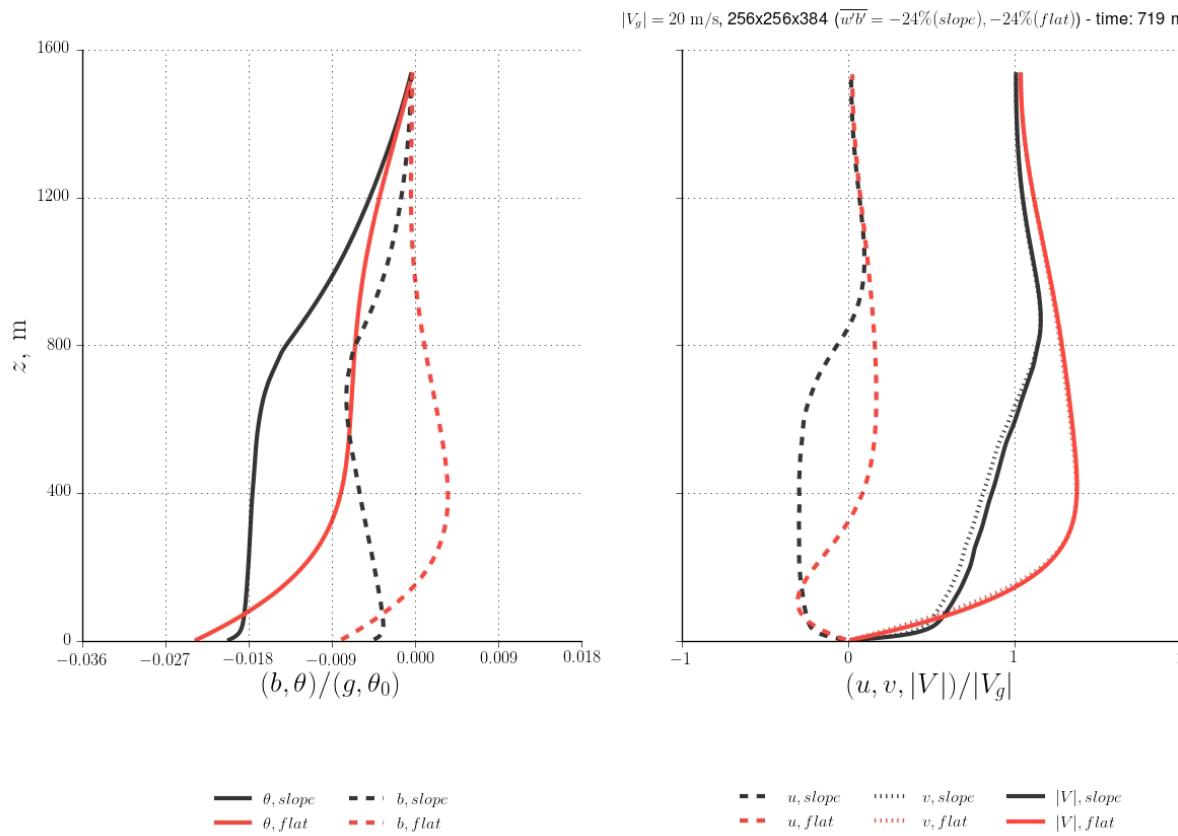


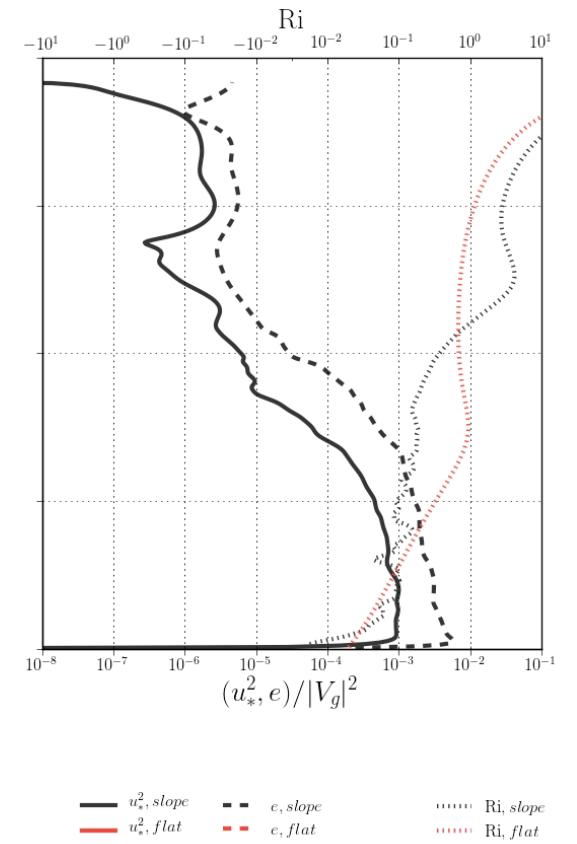
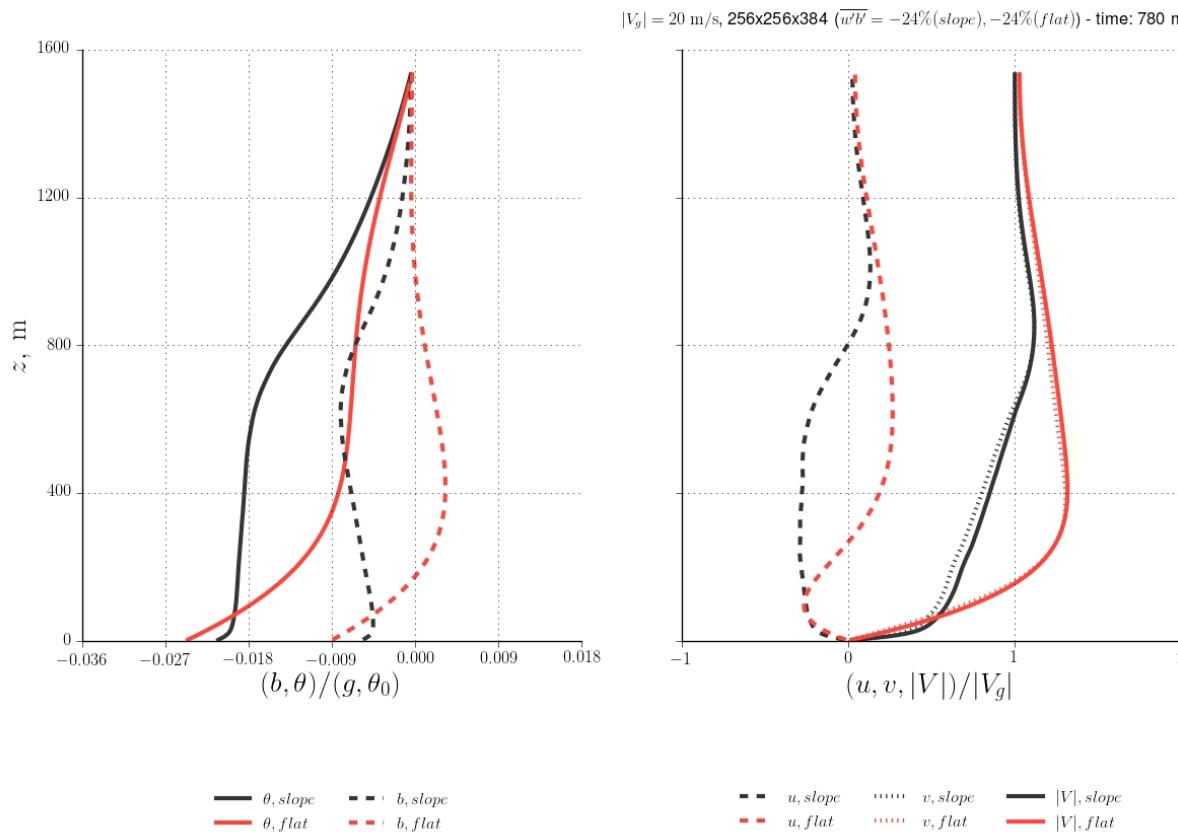


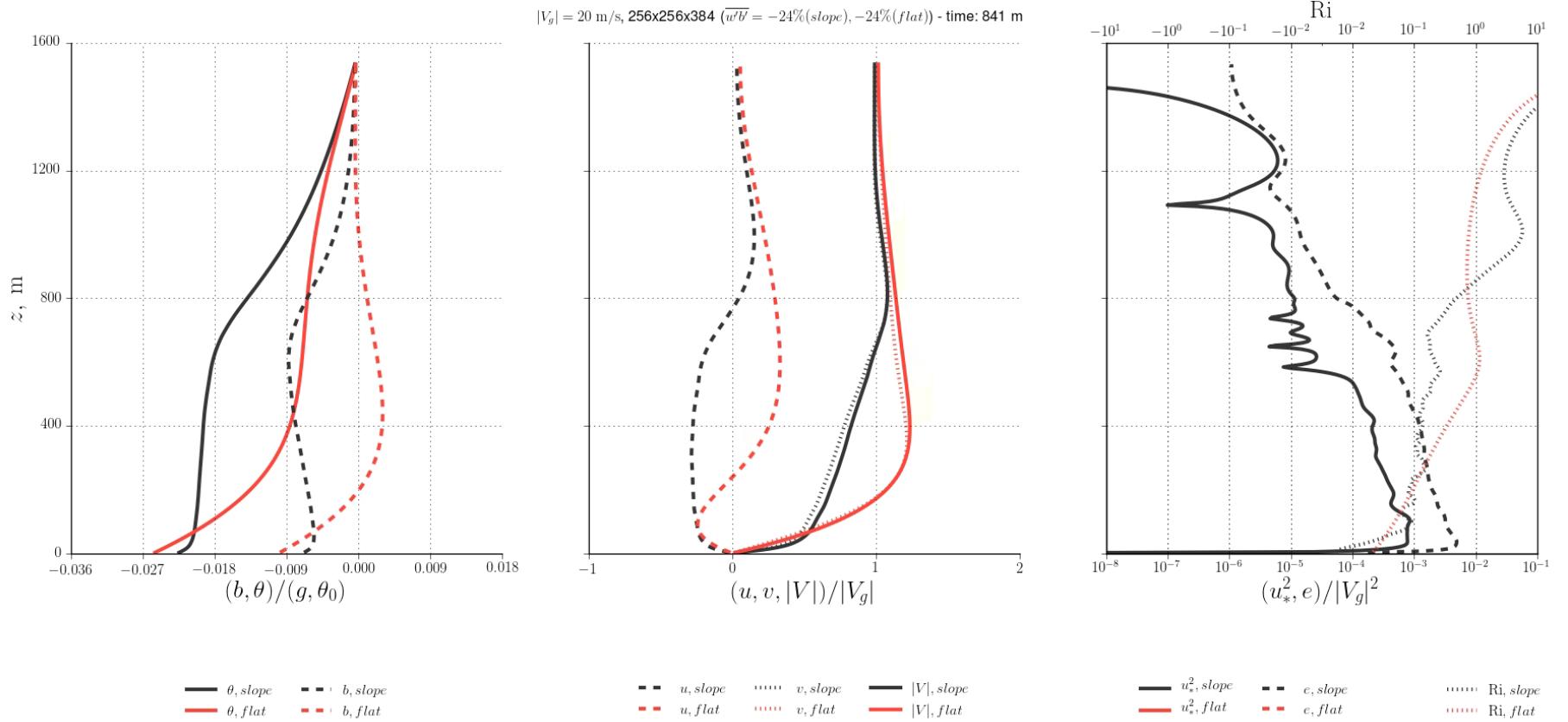


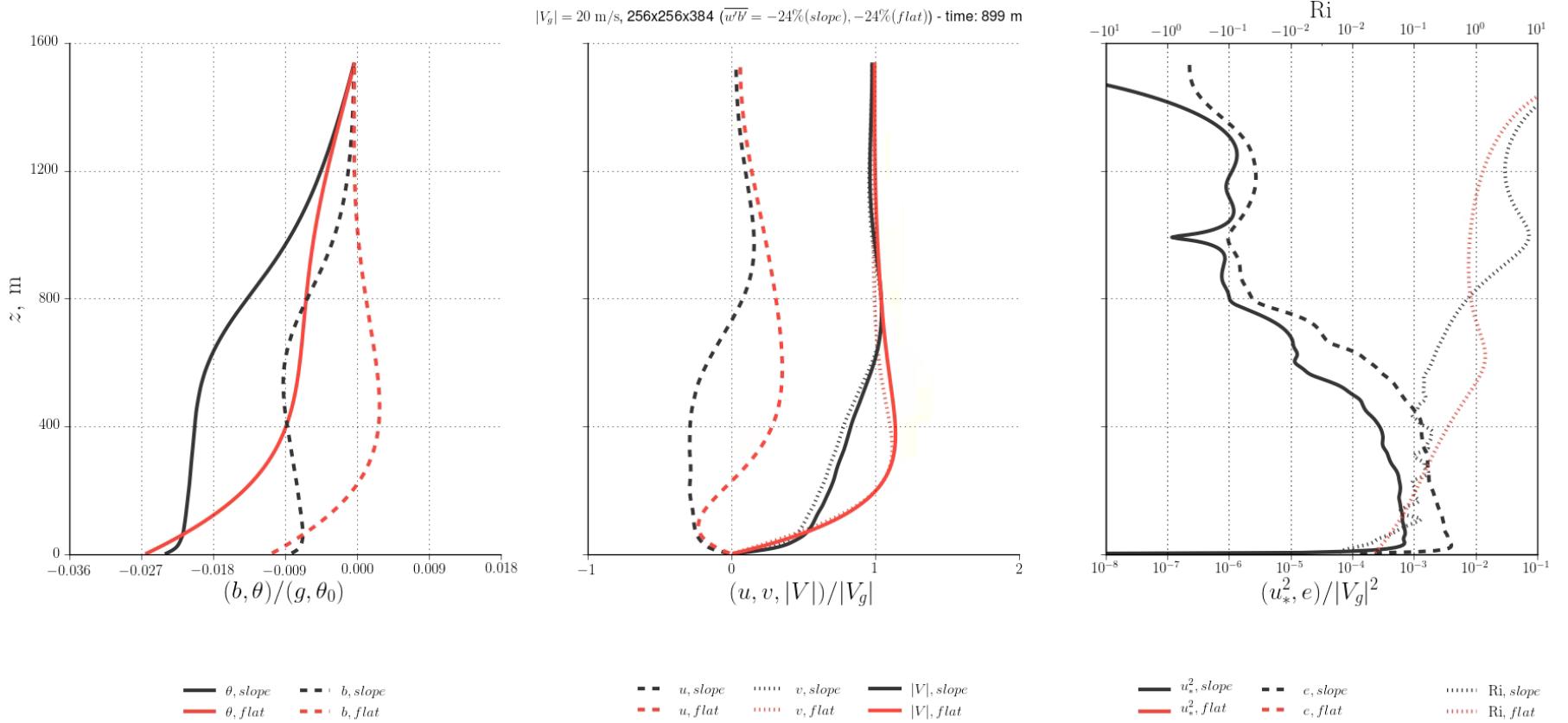






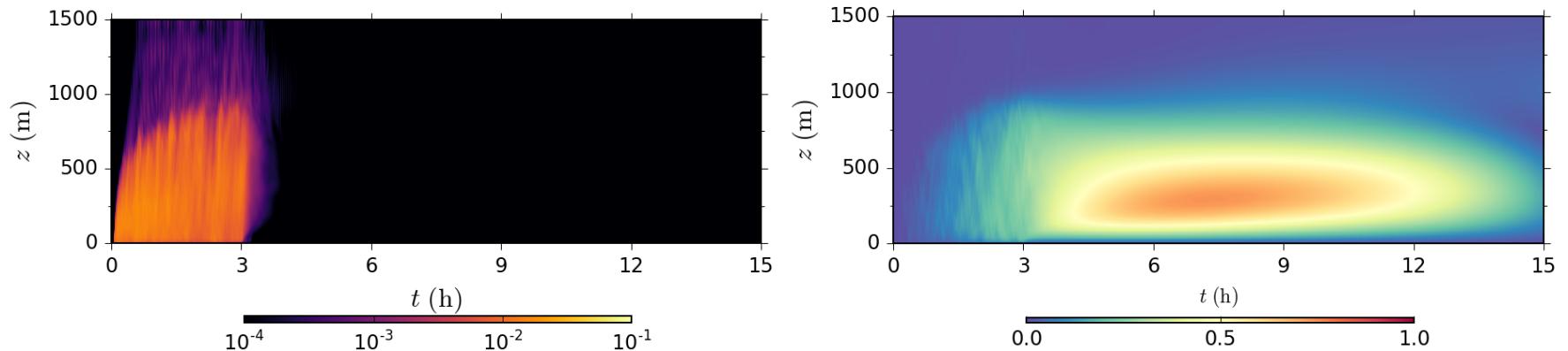
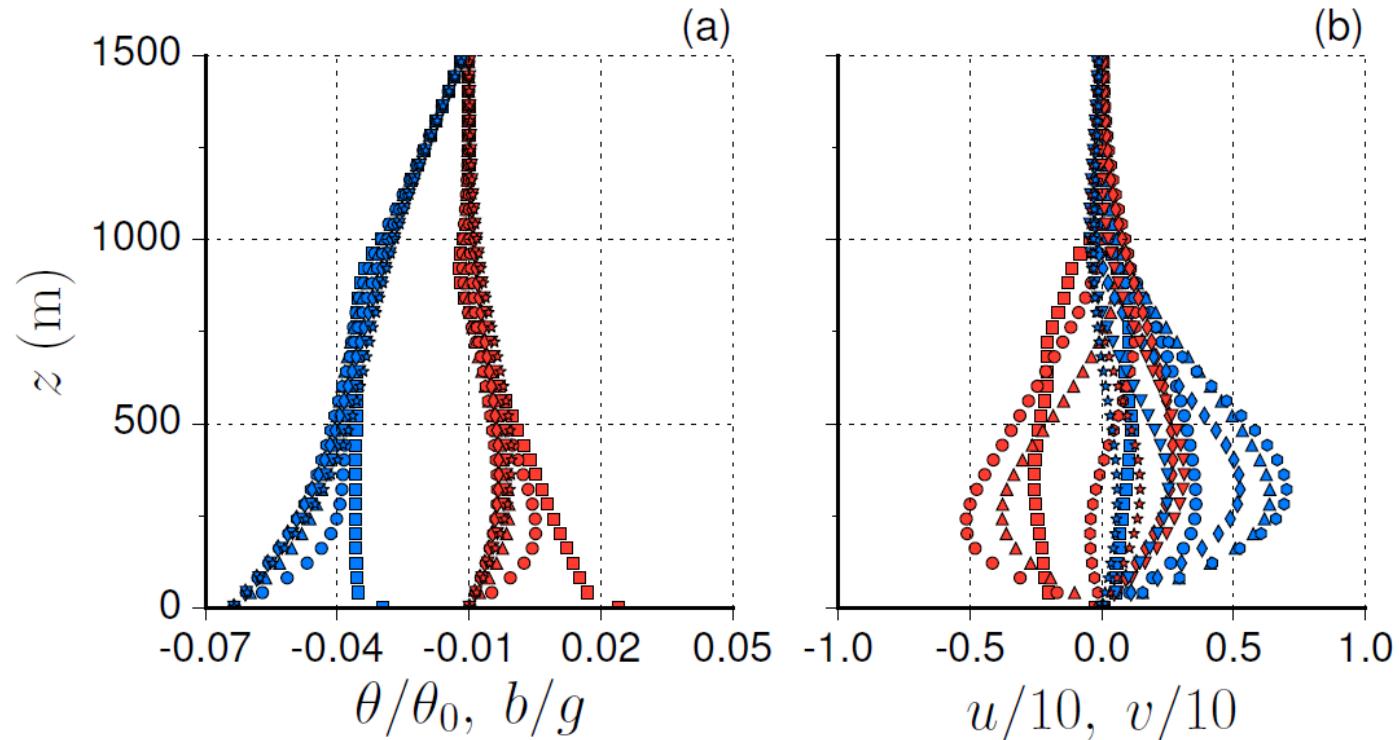






# No external geostrophic forcing is needed for LLJ on a slope!

**Setup:**  $\alpha = 0.18^\circ$ ,  $V_g = 0$ ,  $b_{\text{sd}} = 0.1 \text{ m s}^{-2}$ ,  $b_{\text{sn}} = 0$ ,  $N = 2 \times 10^{-2} \text{ s}^{-1}$



# Conclusions

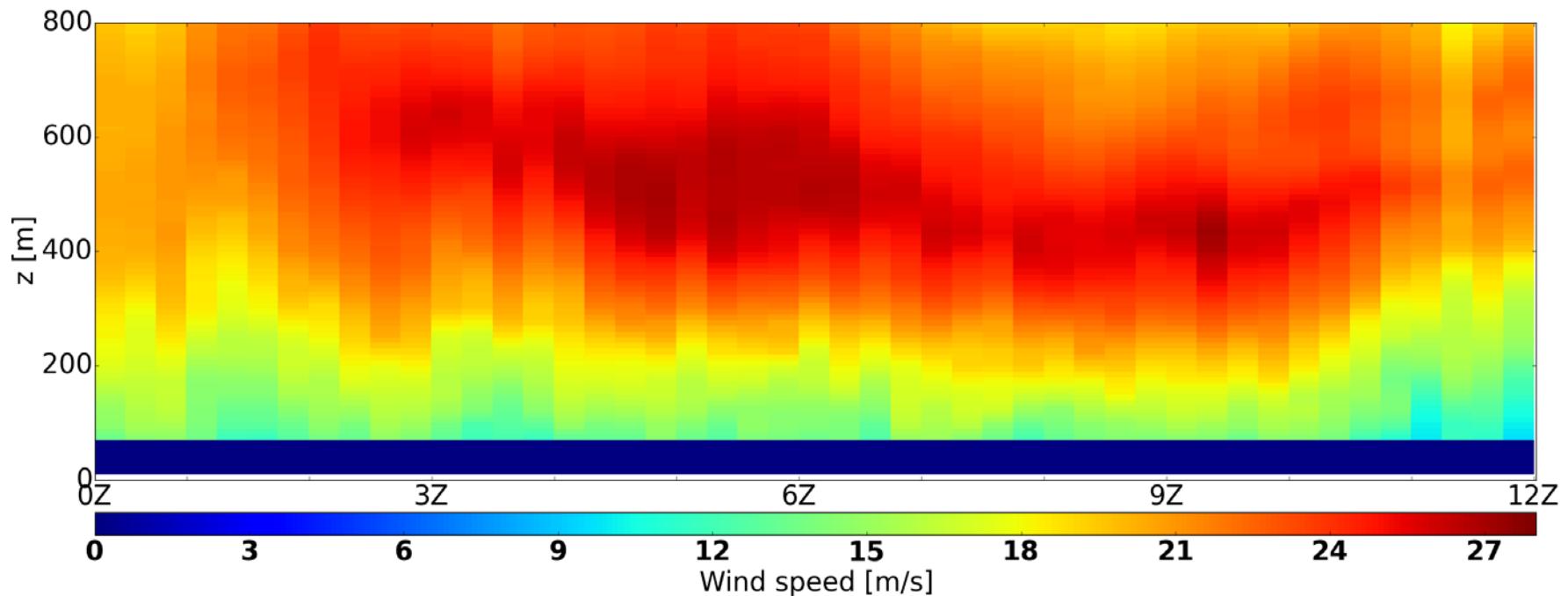
- Shallow slope affects the LLJ structure and evolution in multiple ways
- Along-slope advection of environmental potential temperature can turn a quiet night into a turbulent one, leading to drastic changes in the LLJ properties
- Over a slope, a pronounced nighttime LLJ can develop in the absence of external geostrophic forcing
- LLJ evolution over a slope depends on the type of surface forcing (buoyancy versus buoyancy flux) much stronger than evolution of the flat-terrain LLJ
- Daytime flow preconditioning plays especially important role in LLJ over a slope

Support from NSF (AGS-1359698) is gratefully acknowledged!

# Observed nocturnal LLJ over Great Plains

October 24, 2012, northern Oklahoma

Wind speed from Doppler lidar scans



# Scaled governing equations

**Scales:**  $V = |V_g|$  velocity;  $H$  (boundary-layer depth  $\sim$  domain height) length;  
 $HV^{-1}$  time;  $V^2H^{-1}$  buoyancy;  $V^2$  pressure.

## Normalized governing equations:

$$\frac{\partial u_n}{\partial t_n} + u_n \frac{\partial u_n}{\partial x_n} + v_n \frac{\partial u_n}{\partial y_n} + w_n \frac{\partial u_n}{\partial z_n} = -\frac{\partial \pi_n}{\partial x_n} + Ro^{-1}(v_n - 1) - b_n \sin \alpha + Re^{-1} \left( \frac{\partial^2 u_n}{\partial x_n^2} + \frac{\partial^2 u_n}{\partial y_n^2} + \frac{\partial^2 u_n}{\partial z_n^2} \right), \quad (6)$$

$$\frac{\partial v_n}{\partial t_n} + u_n \frac{\partial v_n}{\partial x_n} + v_n \frac{\partial v_n}{\partial y} + w_n \frac{\partial v_n}{\partial z_n} = -\frac{\partial \pi_n}{\partial y_n} - Ro^{-1}u_n + Re^{-1} \left( \frac{\partial^2 v_n}{\partial x_n^2} + \frac{\partial^2 v_n}{\partial y_n^2} + \frac{\partial^2 v_n}{\partial z_n^2} \right), \quad (7)$$

$$\frac{\partial w_n}{\partial t_n} + u_n \frac{\partial w_n}{\partial x_n} + v_n \frac{\partial w_n}{\partial y_n} + w_n \frac{\partial w_n}{\partial z_n} = -\frac{\partial \pi_n}{\partial z_n} + b_n \cos \alpha + Re^{-1} \left( \frac{\partial^2 w_n}{\partial x_n^2} + \frac{\partial^2 w_n}{\partial y_n^2} + \frac{\partial^2 w_n}{\partial z_n^2} \right), \quad (8)$$

$$\frac{\partial b_n}{\partial t_n} + u_n \frac{\partial b_n}{\partial x_n} + v_n \frac{\partial b_n}{\partial y_n} + w_n \frac{\partial b_n}{\partial z_n} = BuRo^{-2}(u_n \sin \alpha - w_n \cos \alpha) + Re^{-1} \left( \frac{\partial^2 b_n}{\partial x_n^2} + \frac{\partial^2 b_n}{\partial y_n^2} + \frac{\partial^2 b_n}{\partial z_n^2} \right), \quad (9)$$

$$\frac{\partial u_n}{\partial x_n} + \frac{\partial v_n}{\partial y_n} + \frac{\partial w_n}{\partial z_n} = 0. \quad (10)$$

# Dimensionless parameters of the scaled problem

**Rossby number**  $\text{Ro} = VH^{-1}f^{-1}$     **Reynolds number**  $\text{Re} = VH\nu^{-1}$

**Burger number**  $\text{Bu} = N^2 f^{-2}$      $b_s$ -based **Ri number**  $\text{Ri} = -b_s HV^{-2}$

$N$ -based **Richardson number**  $\text{BuRo}^{-2} = N^2 H^2 V^{-2}$

## Atmosphere (ATM):

$$V = 10 \text{ m s}^{-1}, \quad H = 10^3 \text{ m}, \quad f = 10^{-4} \text{ s}^{-1}, \quad N = 10^{-2} \text{ s}^{-1}, \quad \nu = 10^{-5} \text{ m}^2 \text{s}^{-1}, \quad |b_s| = 0.1 \text{ m s}^{-2};$$

$$\text{Ro} = 10^2, \quad \text{Re} = 10^9, \quad \text{Bu} = 10^4, \quad |\text{Ri}| = 1$$

**Length scale range:**  $\eta \sim (\nu^3 V^{-3} H)^{1/4} \sim 10^{-4} \text{ m}$      $\leftrightarrow$      $\sim H = 10^3 \text{ m}$

## Downscaled laboratory analog of atmospheric flow (LEX):

$$V = 0.1 \text{ m s}^{-1}, \quad H = 0.1 \text{ m}, \quad f = 10^{-2} \text{ s}^{-1}, \quad N = 1 \text{ s}^{-1}, \quad \nu = 10^{-6} \text{ m}^2 \text{s}^{-1}, \quad |b_s| = 0.1 \text{ m s}^{-2};$$

$$\text{Ro} = 10^2, \quad \text{Re} = 10^4, \quad \text{Bu} = 10^4, \quad |\text{Ri}| = 1$$

**Length scale range:**  $\eta \sim (\nu^3 V^{-3} H)^{1/4} \sim 10^{-4} \text{ m}$      $\leftrightarrow$      $\sim H = 0.1 \text{ m}$

## Viscous atmosphere (VAT):

$$V = 10 \text{ m s}^{-1}, \quad H = 10^3 \text{ m}, \quad f = 10^{-4} \text{ s}^{-1}, \quad N = 10^{-2} \text{ s}^{-1}, \quad \nu = 1 \text{ m}^2 \text{s}^{-1}, \quad |b_s| = 0.1 \text{ m s}^{-2};$$

$$\text{Ro} = 10^2, \quad \text{Re} = 10^4, \quad \text{Bu} = 10^4, \quad |\text{Ri}| = 1$$

**Length scale range:**  $\eta \sim (\nu^3 V^{-3} H)^{1/4} \sim 2 \text{ m}$   $\leftrightarrow$   $\sim H = 10^3 \text{ m}$

## Interpretations:

Integral flow scales in **VAT** are the same as in **ATM**, but the working fluid is way more viscous than the air.

For given  $\alpha$ , the equality of flow numbers (Ro, Re, Bu, and Ri) for the small-scale **LEX** flow and for the large-scale **VAT** flow signifies the similarity of these flows in terms of equality of their dimensionless solutions.

**VAT** simulation may be conceptually interpreted as large-eddy simulation (LES) with constant subgrid eddy diffusivities for momentum and heat/buoyancy.